

EXPLORING HABITABILITY CONDITIONS IN TITAN'S IMPACT RECORD: THE FORMATION OF MENRVA CRATER





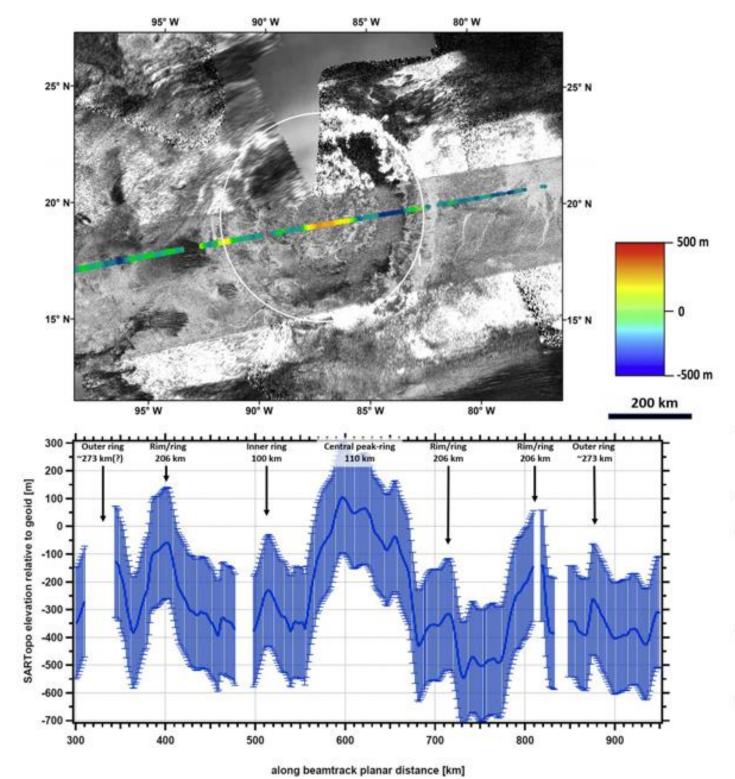
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MENRVA CRATER: THE LARGEST IMPACT ON TITAN

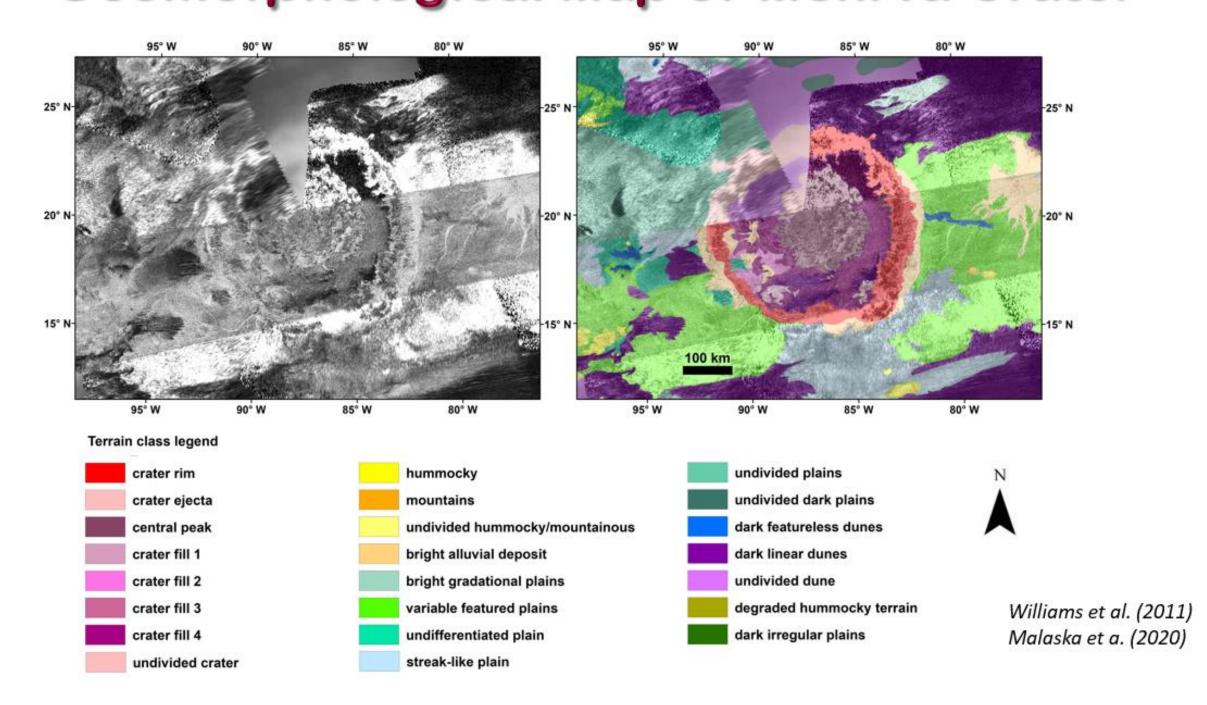
- Among Titan's known impact-related landforms, Menrva crater, at 425 km diameter, stands out due to its large dimensions, making it at least <u>three times larger than the</u> <u>next largest crater</u> (Forseti, ca. 145 km diameter).
- Menrva is a <u>complex impact basin</u>, exhibiting a <u>peak-ring-like structure</u> in its central area.
- Despite degradation, <u>Menrva still exhibits characteristic morpho-structural</u>
 <u>features</u>: rim, two inner annular rings, crater fill, and possible peak-ring central elevations and partially preserved ejecta deposits.

MENRVA CRATER: REMOTE SENSING ANALYSIS

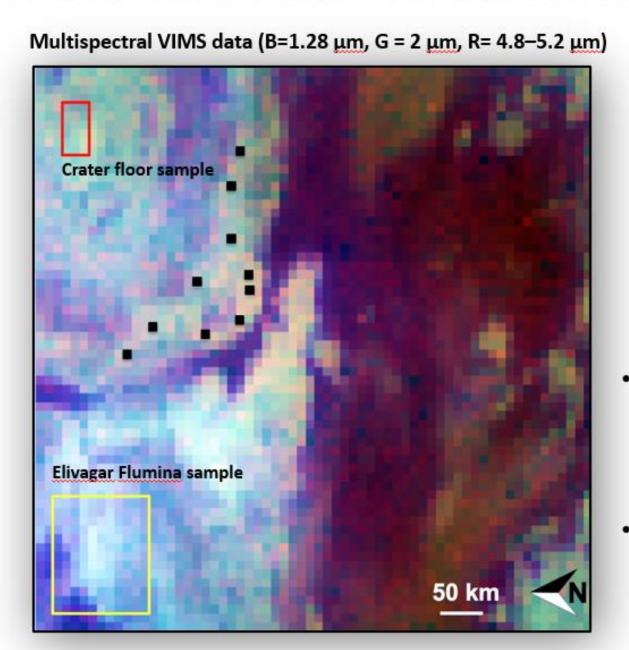


- <u>Topographic transect</u> cutting across the center of the crater using <u>SARTopo</u> data) overlaid on the SAR image.
- Warmer colors represent higher elevations, whereas cooler color the lower ones.
- The relative heights range from ca. <u>-600 to</u>
 <u>+200 m</u> along the transect.
- <u>Topographic section</u> of Menrva Crater along WSW-ENE using <u>SARTopo</u>.
- Error bars are positioned at +/- 200 m.
 The main rim is at ~206 km radius (~412 km diameter) and the inner ring is at ~100 km radius (~200 km diameter).
- The central section (<u>peak-ring</u>) extends out from the center to <u>~55 km</u> (~110 km diameter).
- The subtle <u>outer ring shows at ~273 km</u>
 <u>radius</u>.

Geomorphological Map of Menrva Crater



MENRVA CRATER: REMOTE SENSING SPECTRAL ANALYSIS



Menrva: spectrally a low emissivity but high albedo target

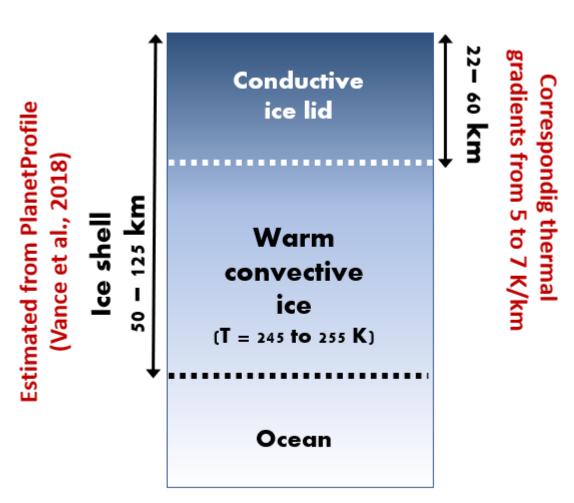
	Menrva		
Unit	Floor	Ejecta	
H ₂ O	18 ± 5%	17 ± 4%	
Tholin	82 ± 19%	83 ± 18% - 85%	
Dark material			
Total organic	85%		
Emissivity	0.90	0.87	
Degradation state	Somewhat d still retains n structura	-	

- The ejecta blanket and crater floor of Menrva seems to comprise a combination of waterice-like spectra, NH₃ ice, CO₂ ice, and tholin components.
- No NH₃ or CO₂ spectra were found, nor the dark material that is present is most of Titan's surface.

Solomonidou et al. (2018) Crósta et al. (2021)

MENRVA CRATER: IMPACT MODELING

- To <u>simulate the formation of a Menrva-like crater</u>, we used <u>iSALE-2D</u>, a multi-material, multi-rheology shock physics code (<u>Melosh</u> et al., 1992; Ivanov et al., 1997; Collins et al., 2004; Wünnemann et al., 2006), which is based on the SALE hydrocode solution algorithm (<u>Amsden</u> et al., 1980).
- Our <u>model inputs are consistent</u> with the earlier modeling study that <u>examined the</u> <u>formation of impact craters on Europa</u> (Silber and Johnson, 2017; 2018), with a few minor adaptations, such as a <u>surface gravity</u> appropriate for Titan.
- Lagrangian tracer particles were implemented to track the material position and state during the crater formation process.



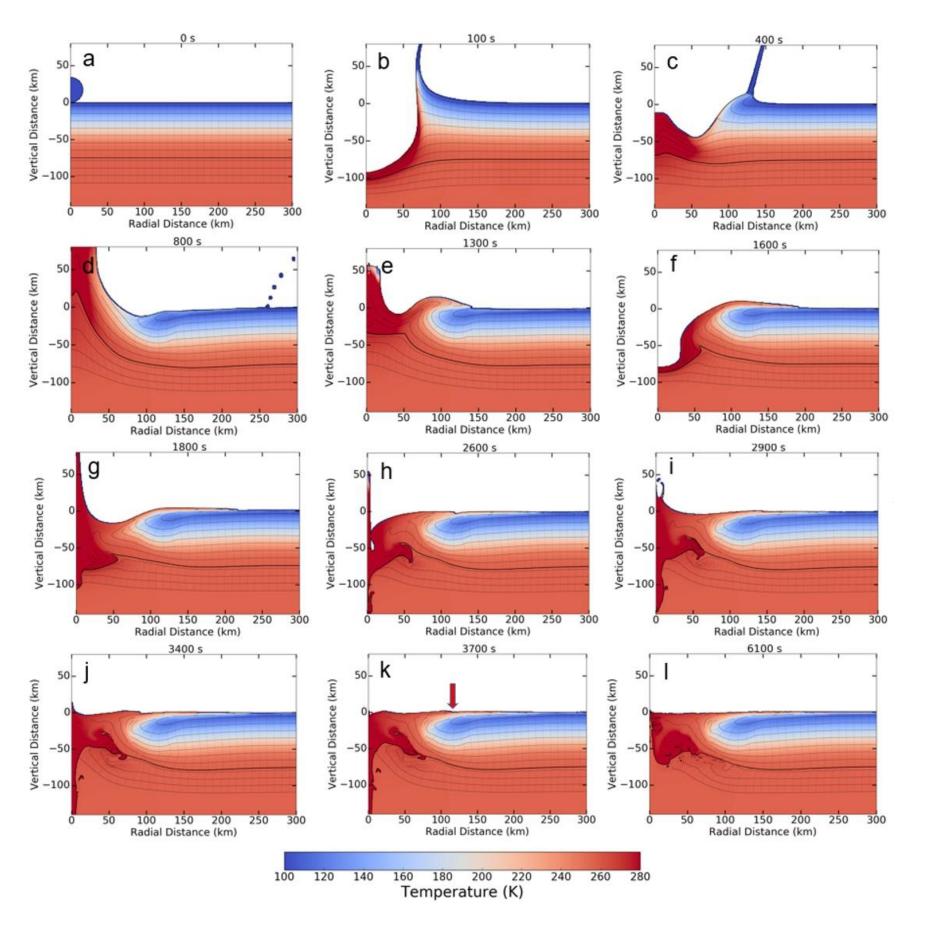
Structure of the icy shell overlying the water ocean and range of parameters used for ice thickness

Description	
Surface temperature (K)	94
Melt temperature at zero pressure (K)	273
Thermal softening parameter	1.2
Cohesion, intact (MPa)	10
Coefficient of internal friction, intact	2
Limiting strength at high pressure, intact (GPa)	0.11
Cohesion, damaged (kPa)	10
Coefficient of internal friction, damaged	0.6
Limiting strength at high pressure, damaged (GPa)	0.11
Equation of state (EOS)	$Tillotson, H_2O$
Impact velocity (km/s)	7
Impact angle (degrees)	45
Impactor diameter (km)	34

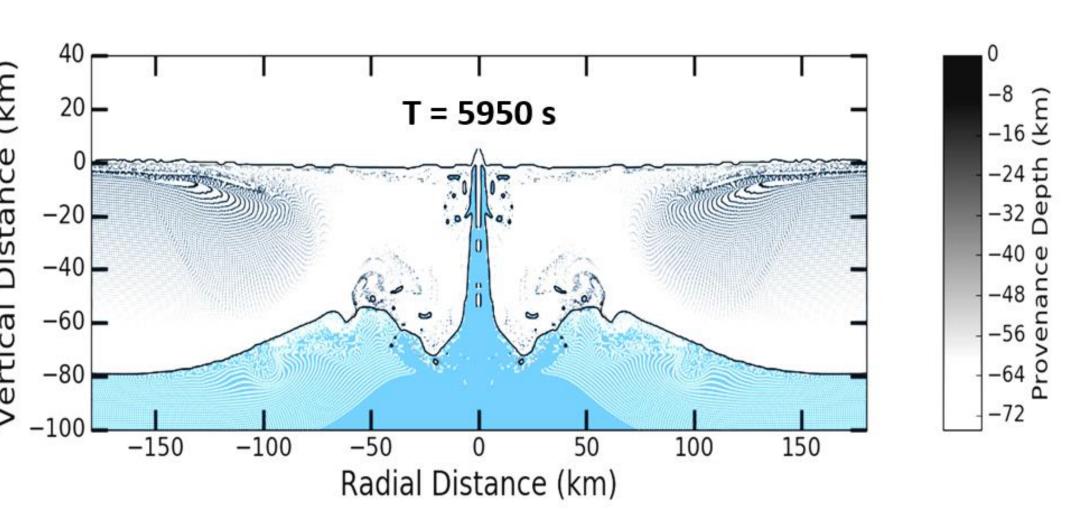
Input parameters for modeling using iSALE-2D

D _{ice}	d _{lid} (km)	T ice (K)	D _i = 34 km	D _i = 28 km
(km)			v _i = 7 km/s	v _i = 10 km/s
50	26	245	Р	Р
50	26	255	Р	Р
75	26	245	Р	Р
75	26	255	Р	Р
75	60	245	Р	Р
75	60	255	Р	Р
75	26	245	Р	Р
100	26	255	F	Р
100	60	245	Р	Р
100	60	255	Р	Р
100	26	245	Р	Р
125	26	255	F	Р
125	60	245	F	Р
125	60	255	Р	F

Parameter space and combinations used for modeling

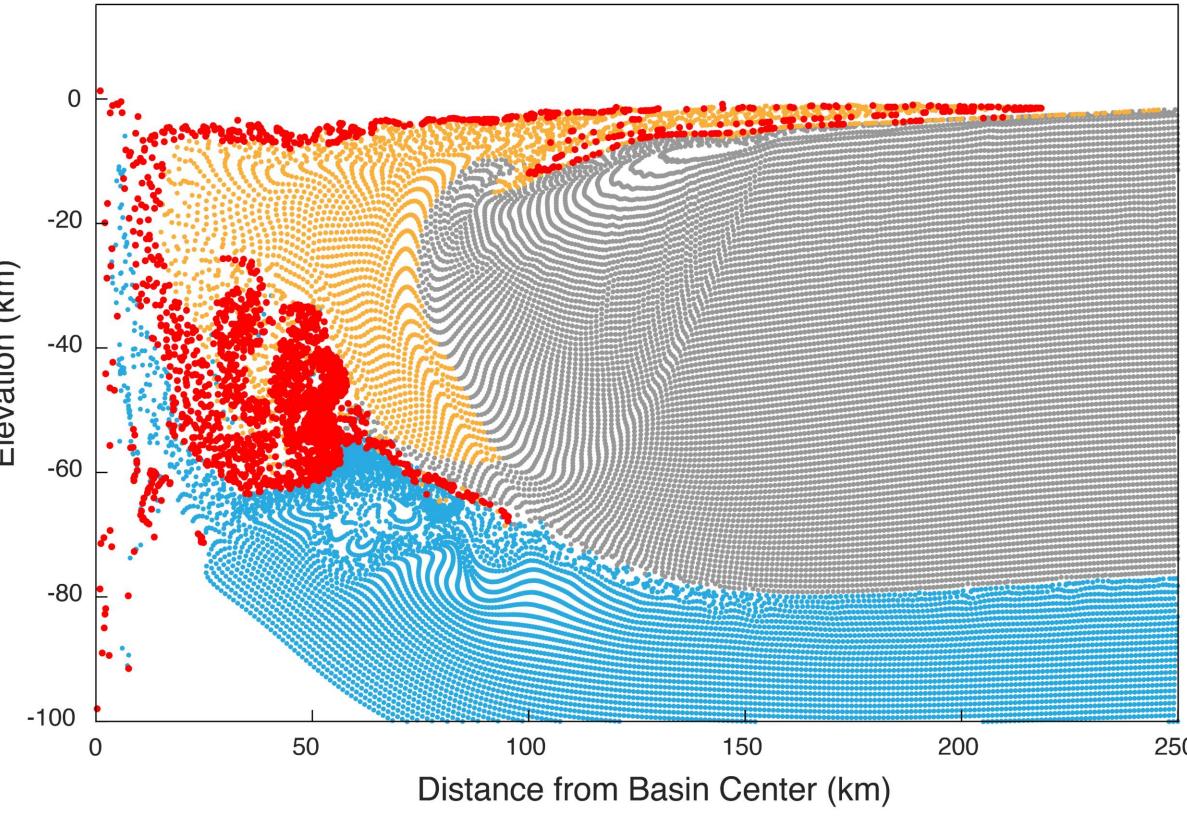


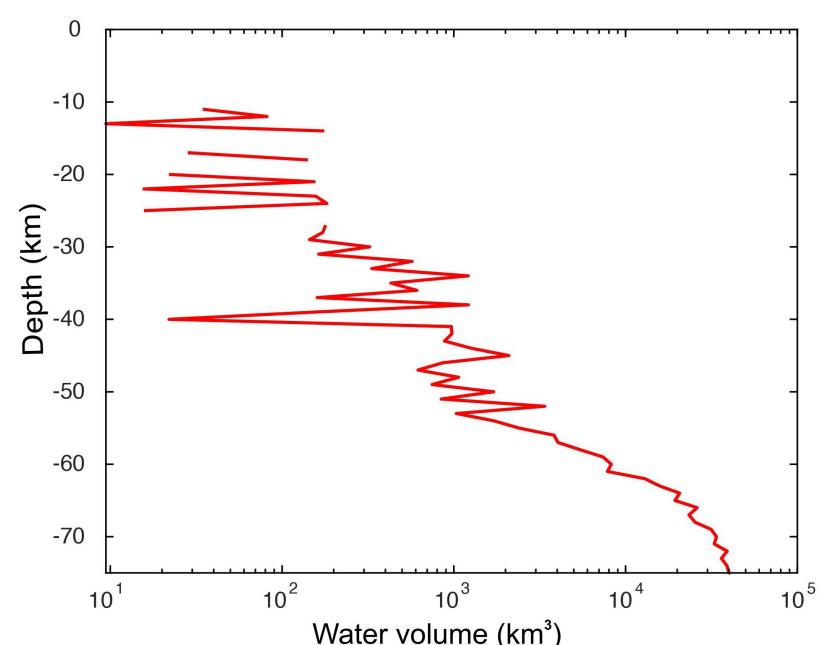
Time series showing the example case of the formation of Menrva. The impactor's diameter is 34 km and the vertical component of impact velocity is 7 km/s. The sequence of figures (7a to 7l) depicts the evolution of the transient crater from the moment of impact (t = 0 s) until it reaches gravitational stabilization (t = 6100 s) (see the text for details on each stage). The red arrow in Fig. 7k indicates the approximate position of the inner ring at ca. 130 km from the center.



A <u>substantial upwelling</u> takes place, with <u>vertical mixing and material</u> <u>exchange between the surface and the ocean</u>, at distances of <u>up to 70 km</u> <u>from the center</u> and <u>up to ca. 80 km in depth</u>. Tracers within the ice shell are colored according to their original depth for the same model shown in the previous figure.

Peak Shock Pressures for: 34km Projectile 75km Lid 60km Conductive Lid 255K Temp 5000s





Total water volumes as a function of depth for the example case

CONCLUSIONS

- Large hypervelocity impacts can have an important role in creating habitable environments or niches. When
 viewed from the perspective of a geobiological process, impact cratering can influence the biological evolution
 of planetary bodies and control planetary habitability.
- Our models clearly indicate that our working hypotheses were basically met in most of the scenarios tested, i.e.,
 (i) Titan's ice shell is indeed breached by a large impact, such as the one that formed Menrva crater and,
 possibly, even smaller impacts depending on the thickness of the icy shell at the time of the event; (ii) materials
 from the three uppermost layers on Titan's interior (organics, ice, and ocean) mix in considerable amounts
 because of deformation processes.
- Deformation processes associated with crater formation provide two of the necessary conditions for a habitable world: the adequate substrates (organic compounds, ice, and water) and large enough volumes of these materials to provide the media for prebiotic life development.
- A third, and key condition, is temperature. Our models showed that relatively large volumes of these materials
 may be heated and melted, mainly at the central area of the crater; as for the duration of these relatively warm
 conditions, inferences can be drawn from Earth's analogues, which show evidence of long-lasting hot fluid
 circulation, of up to a million years.

Acknowledgments

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