

Downscaling Global Warming with a Regional Ocean- Atmosphere Model over the Tropical Atlantic

Role of equatorial ocean dynamics:
equatorial upwelling and ocean mesoscale variability

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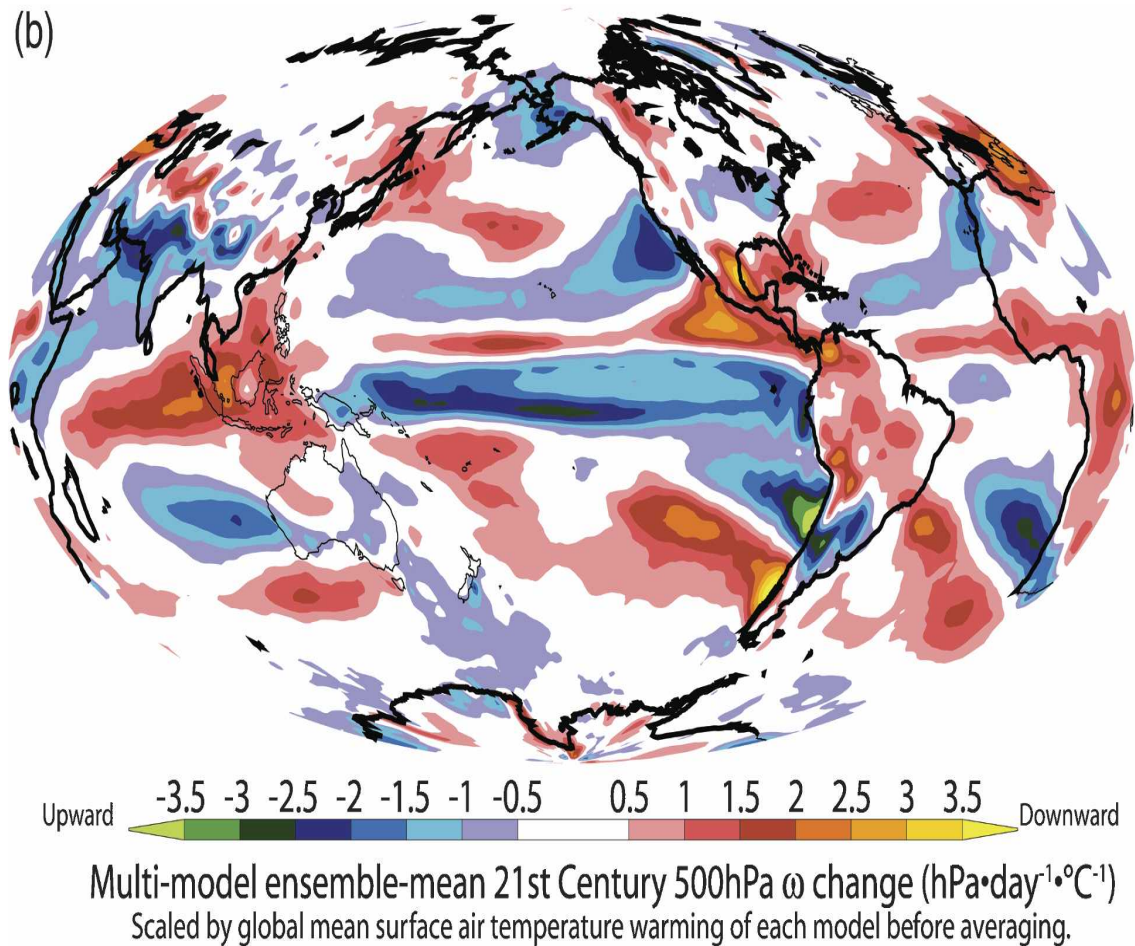
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Also thanks to
Raghu Murtugudde, Markus Jochum,
and Art Miller

Introduction:

Weakening of Walker circulation and ocean heat transport

Multi-model ensemble change
(A1B-20C) in ω (500hPa)

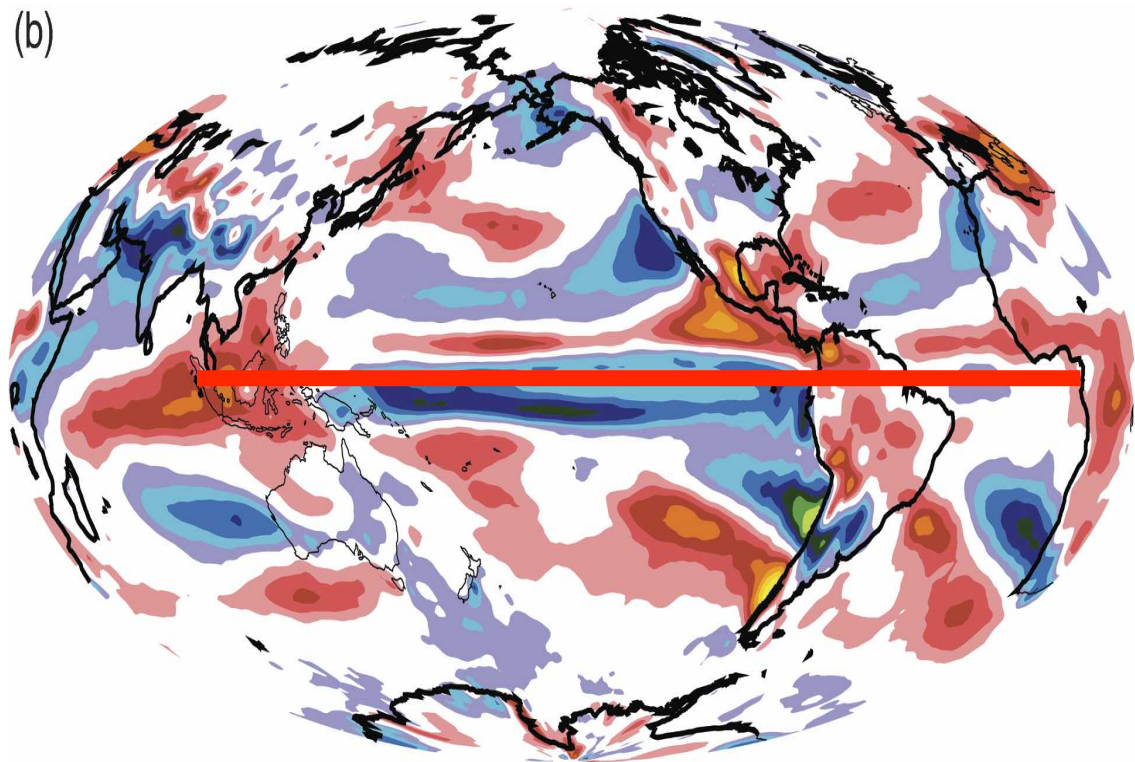


Vecchi and Soden 2007

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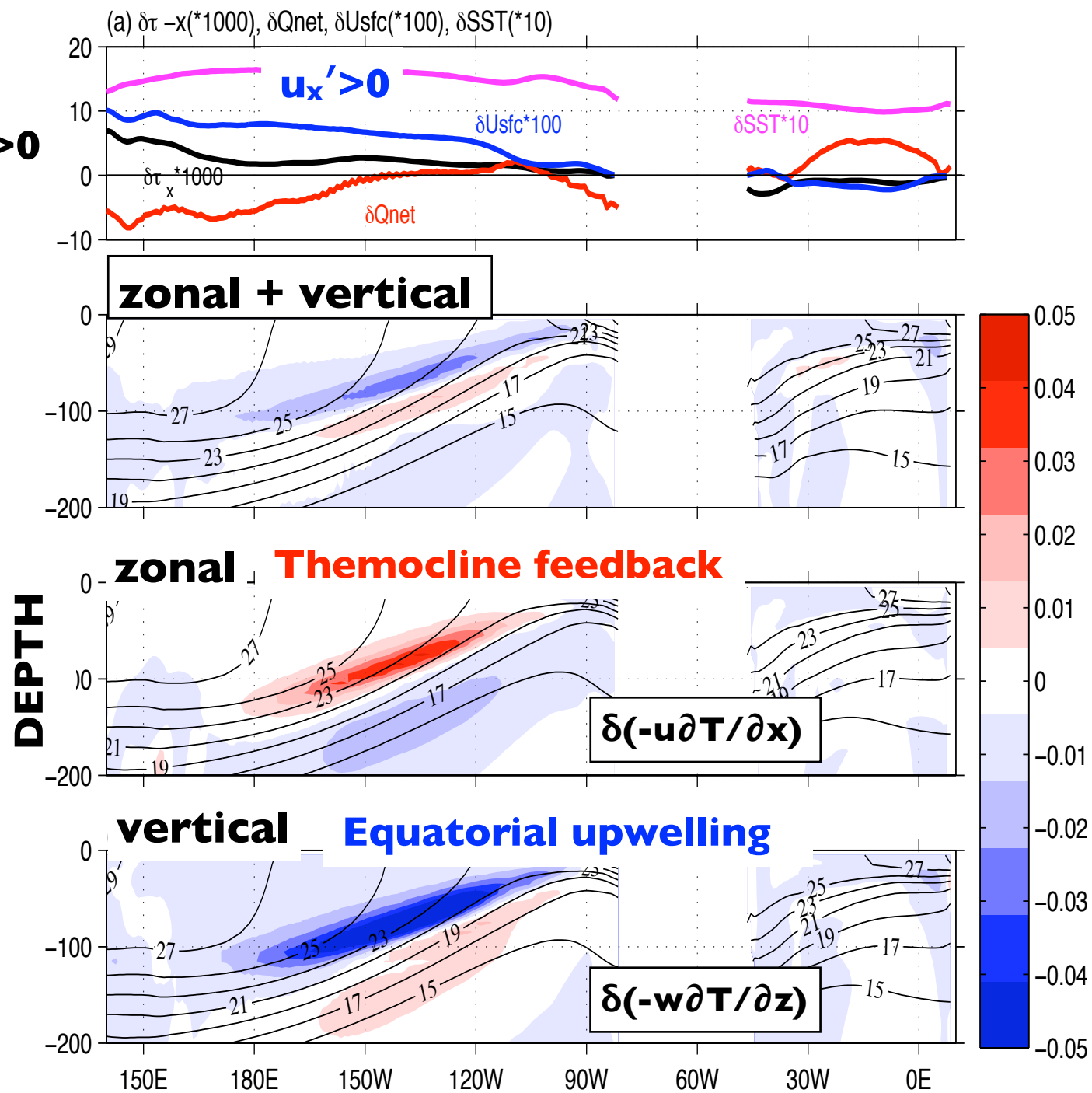


$\tau_x' > 0$

Upward -3.5 -3 -2.5 -2 -1.5 -1 -0.5 0.5 1 1.5 2 2.5 3 3.5 Downward
Multi-model ensemble-mean 21st Century 500hPa ω change ($\text{hPa}\cdot\text{day}^{-1}\cdot\text{C}^{-1}$)
Scaled by global mean surface air temperature warming of each model before averaging.

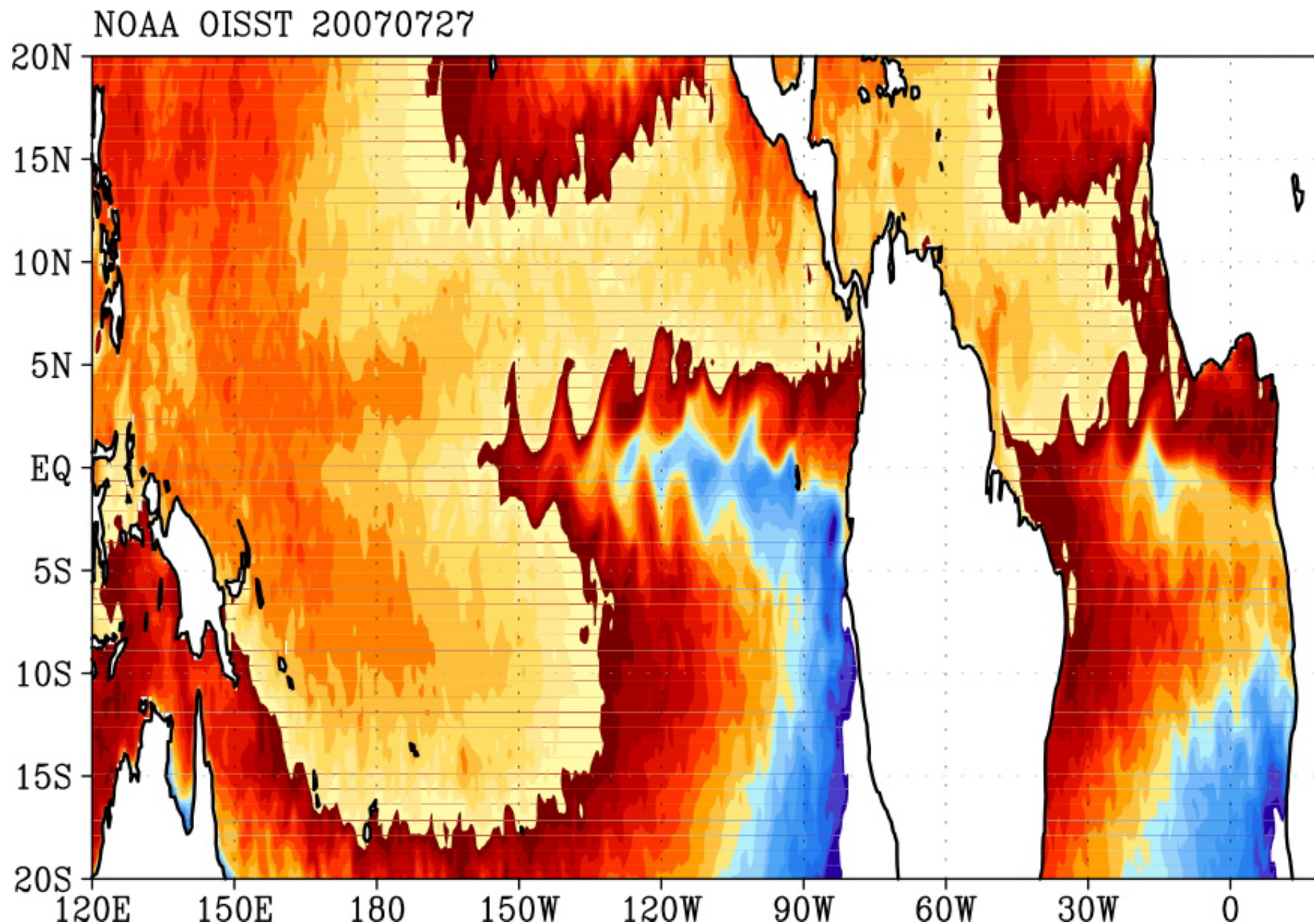
Vecchi and Soden 2007

Change in ocean heat transport



GFDL CM2.1 10-member ensemble
(2046-2050) - (1996-2000)

- **Tropical Instability Waves (TIWs)** are the undulations of equatorial SST front in the Pacific and Atlantic.
- Generated by **oceanic intrinsic instability**.
- **Primarily sub-seasonal, but important for low-frequency tropical climate**.
- Not well-resolved in the IPCC-AR4 models. So we need to **downscale**.

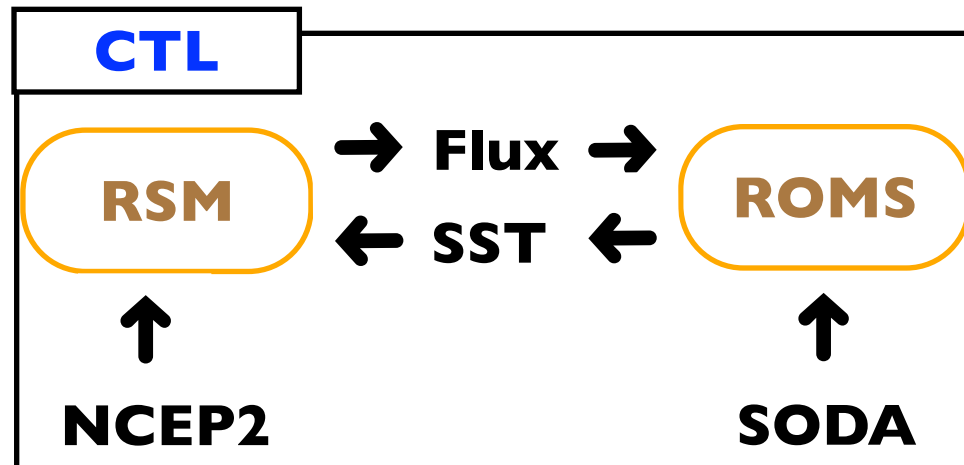


Model and Experiments

Scripps Coupled Ocean-Atmosphere Regional Model*

Atmosphere: Regional Spectral Model (Scripps RSM)

Ocean: Regional Ocean Modeling System (ROMS)



- **CTL**: RSM (NCEP2 6hrly) + ROMS (SODA monthly)
- 25 km ROMS + 50 km RSM
- Daily coupling
- 28-yr. integration: 1980-2007

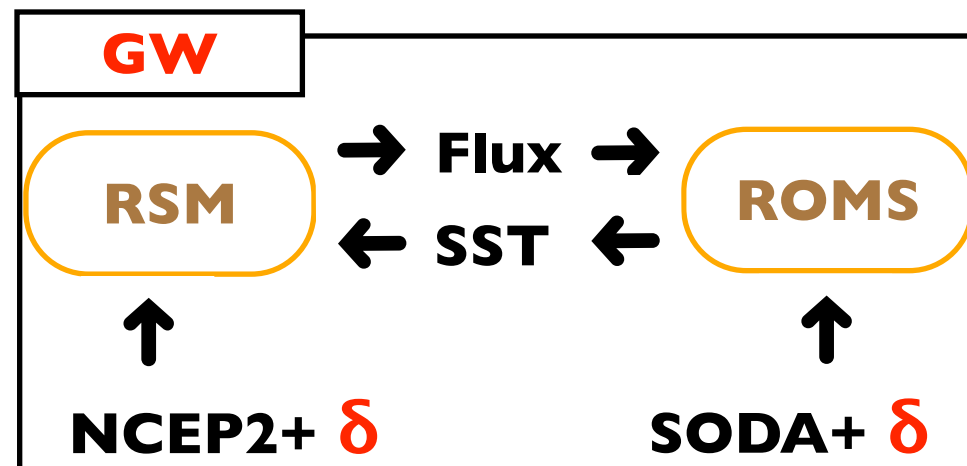
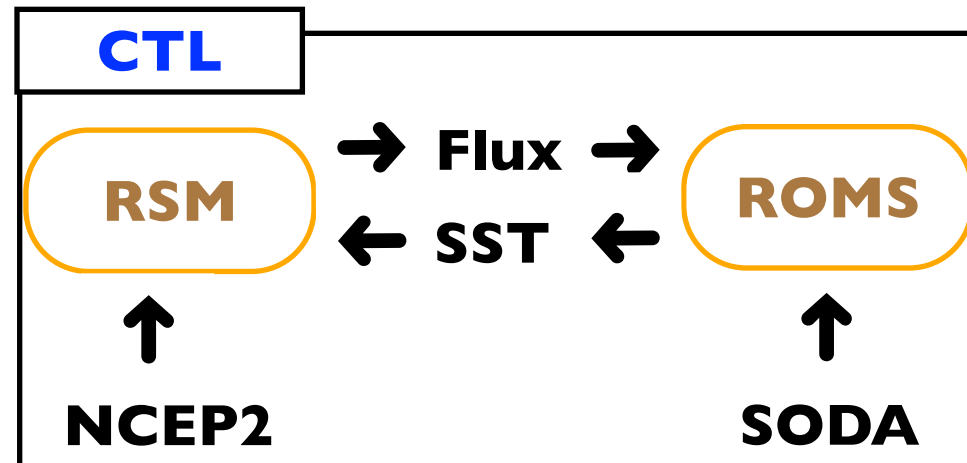
***Seo, Miller and Roads, 2007**: The Scripps Coupled Ocean-Atmosphere Regional (SCOAR) model, with applications in the eastern Pacific sector. *Journal of Climate*

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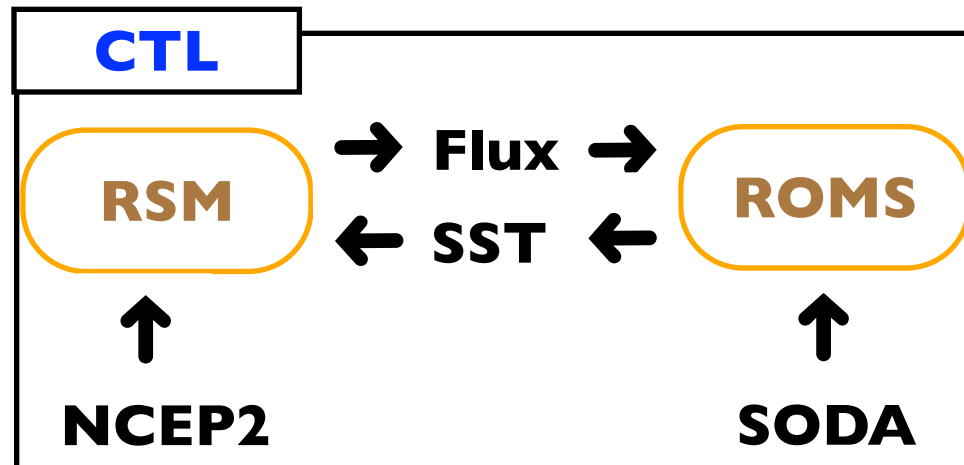
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(2045-2050: A1B) - (1996-2000: 20C)
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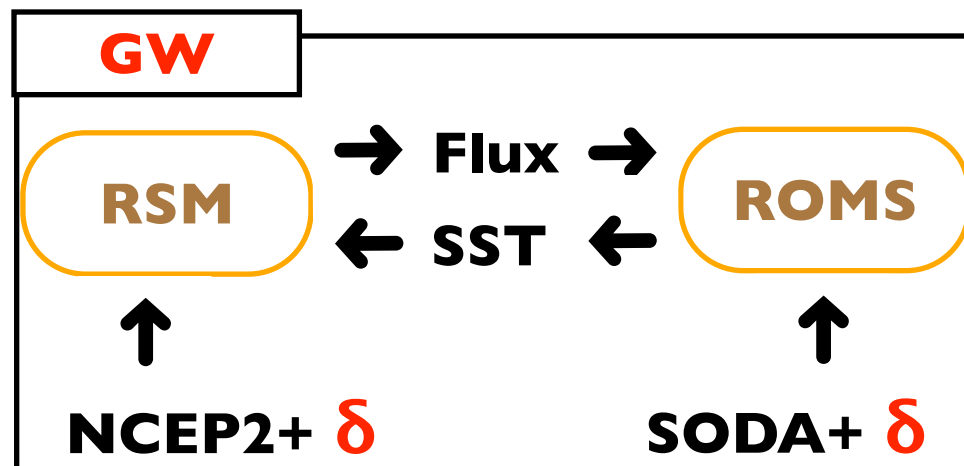
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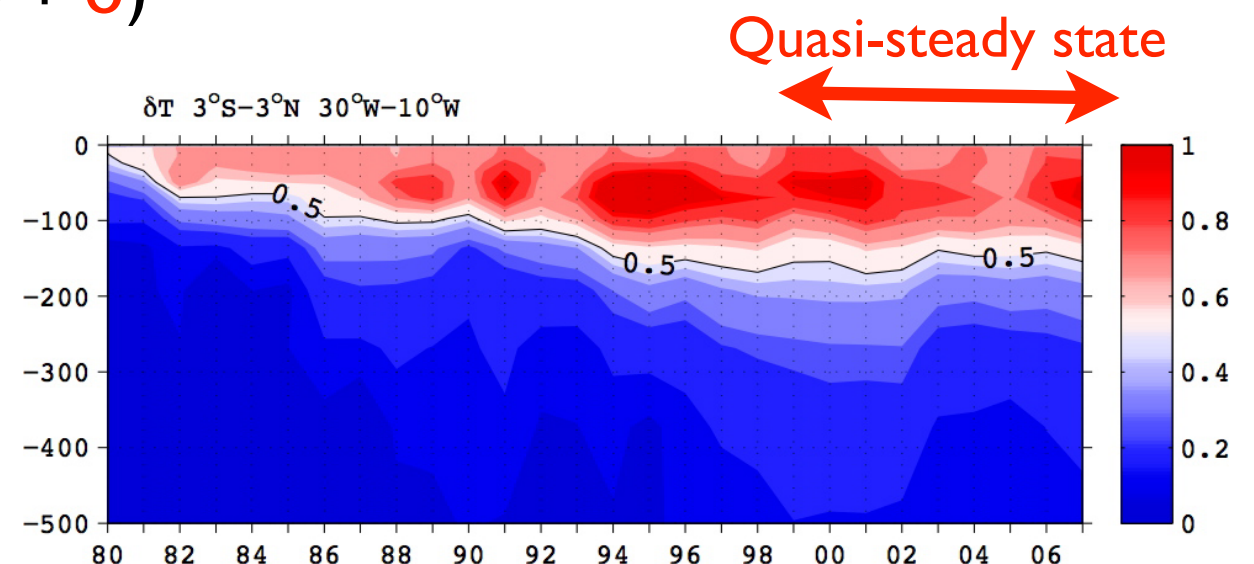


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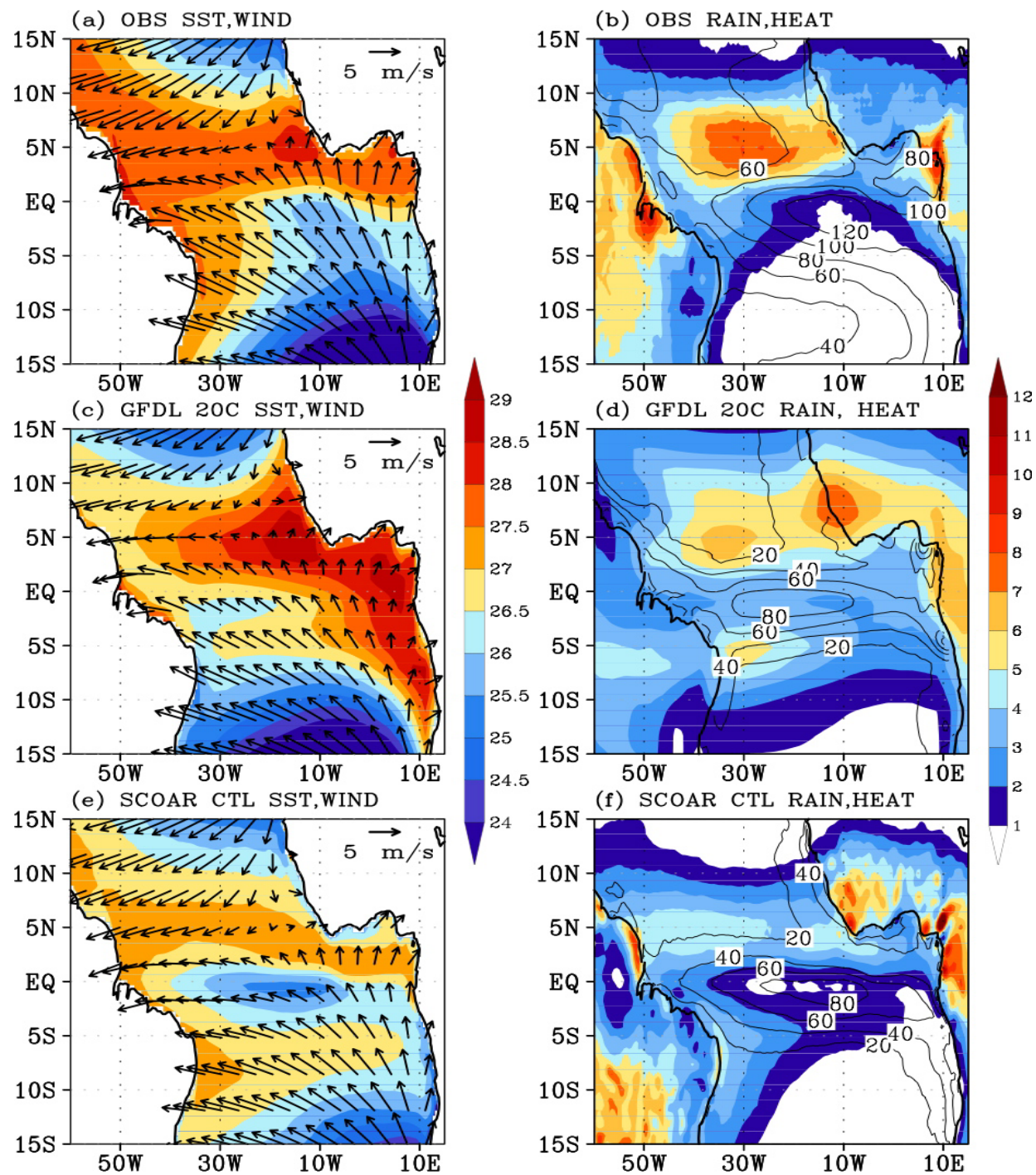
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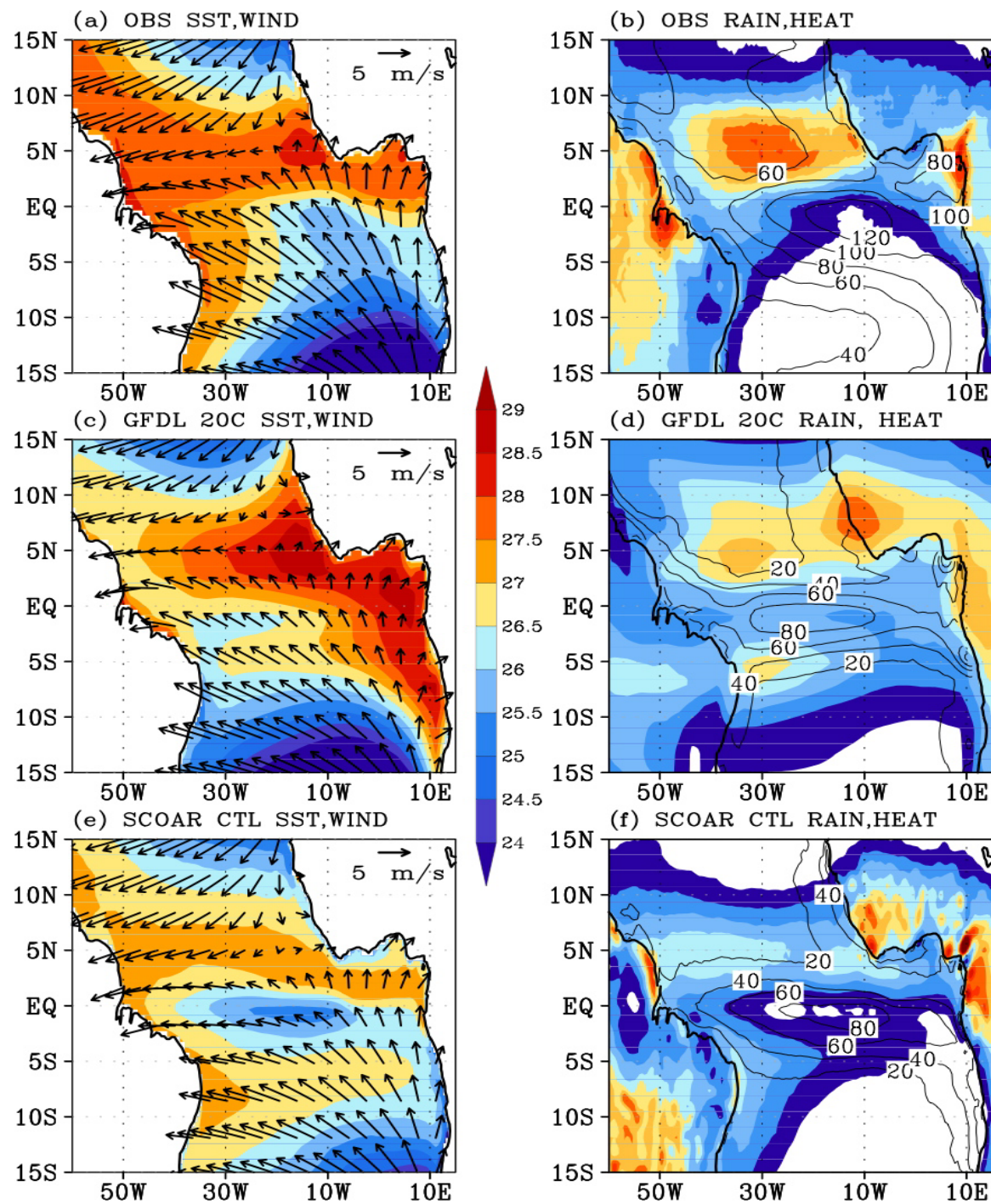
Simulation of present-day climate and global warming response:
Annual mean SST, surface winds, and precip.

Simulation of present-day climate



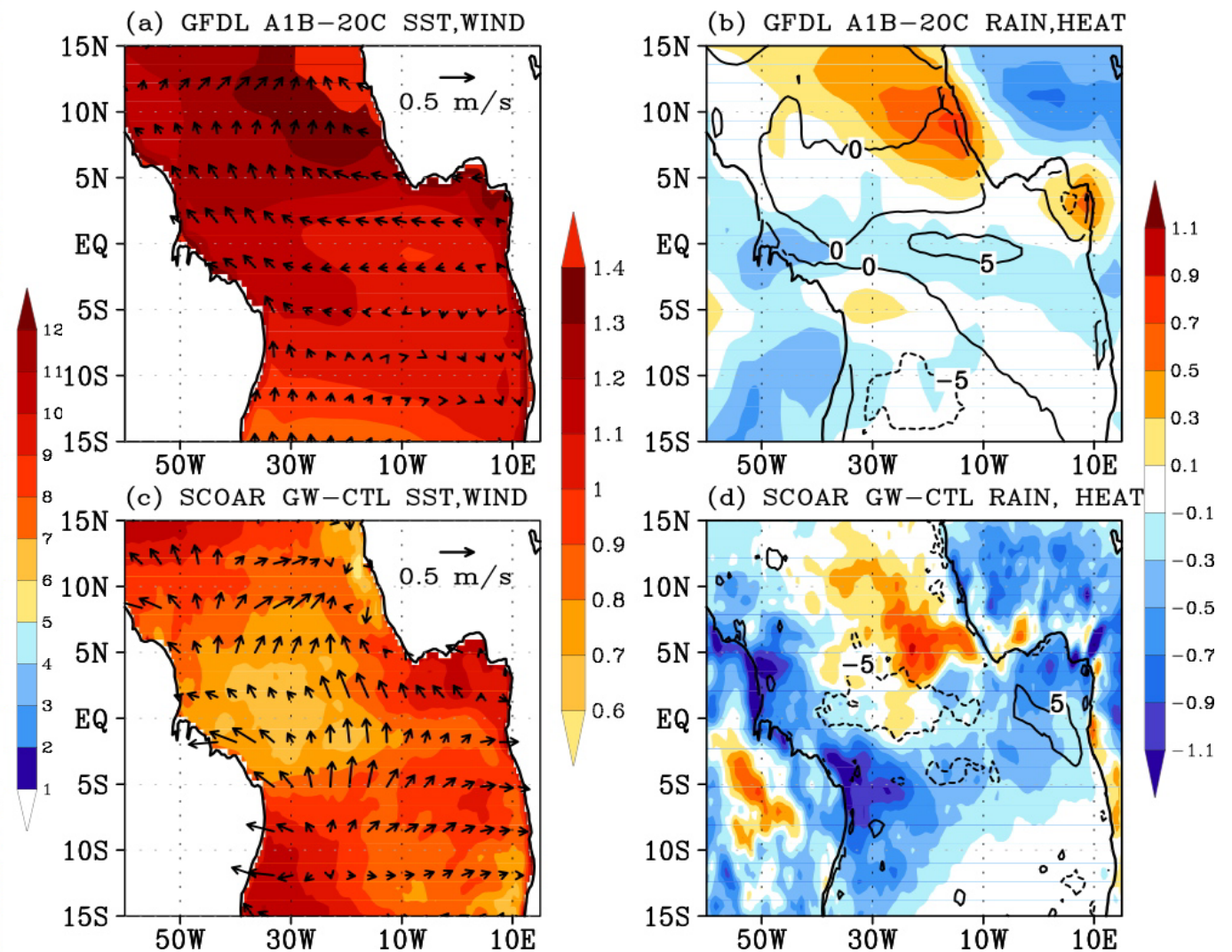
- Zonal SST gradient and equatorial cold tongue in SCOAR

Simulation of present-day climate



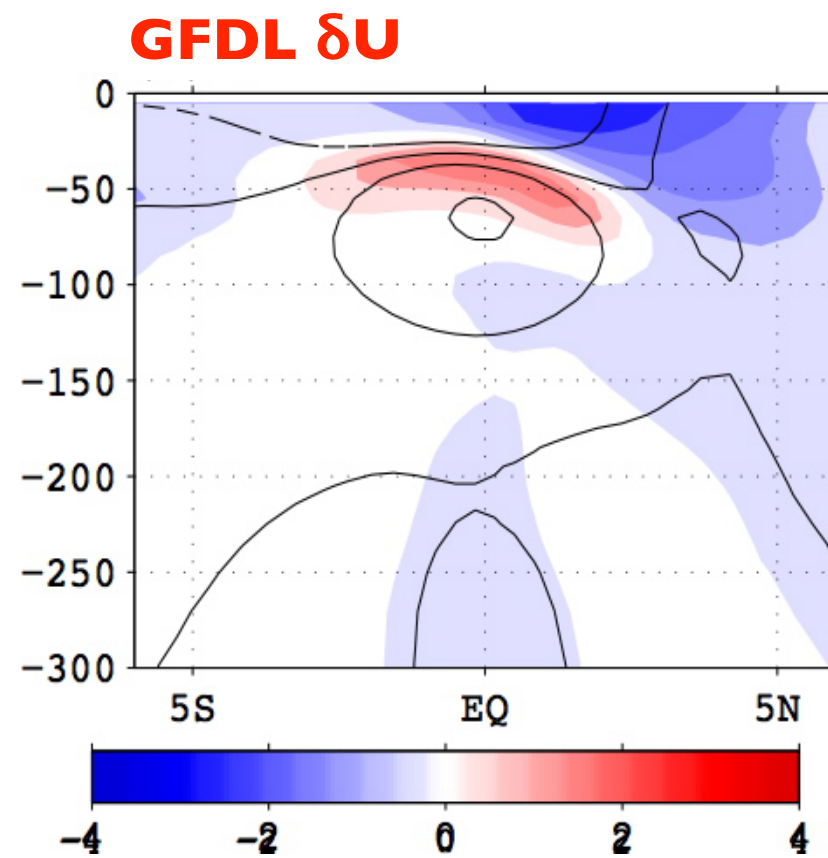
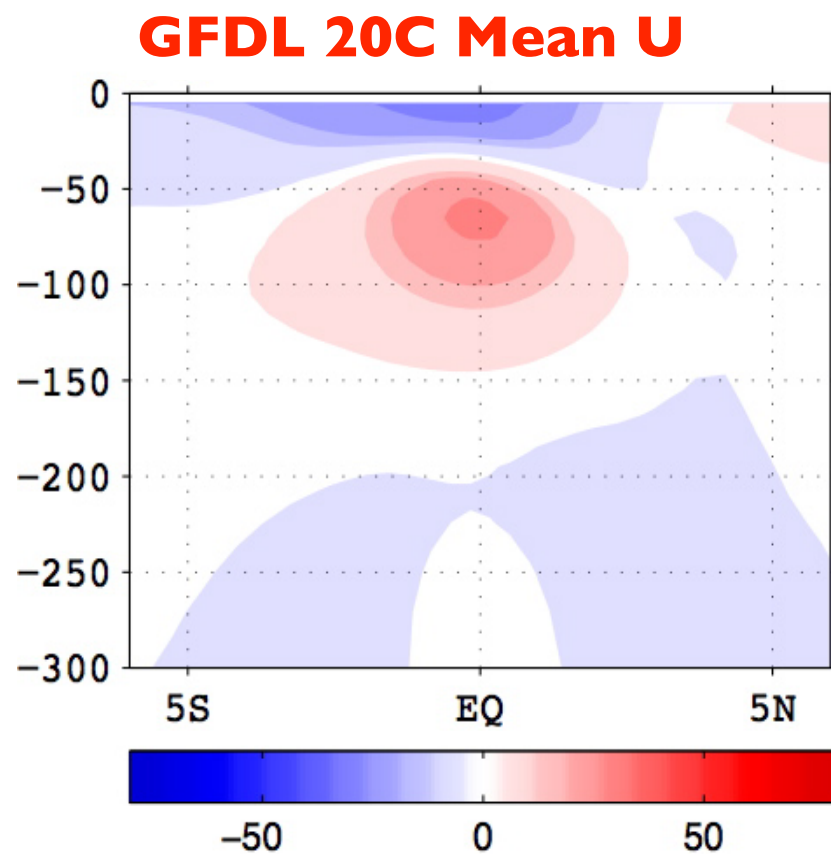
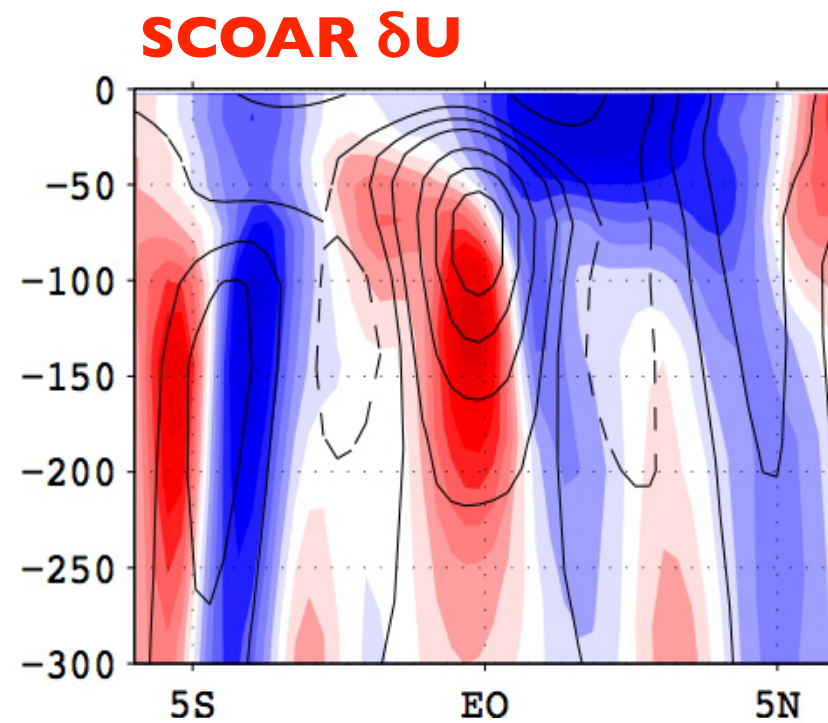
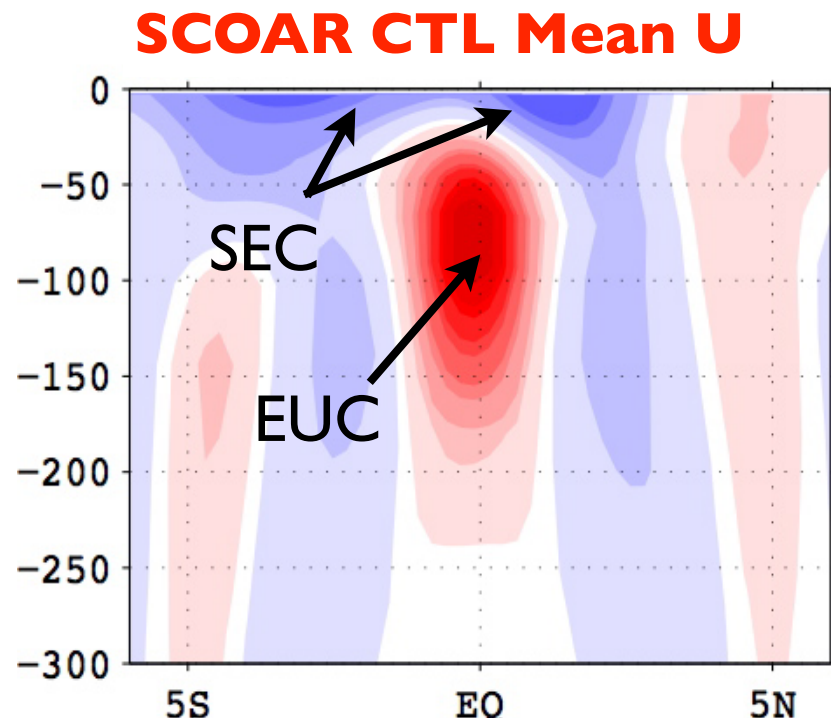
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GW response



- Reduced warming in the equator
- Intensified cross-equatorial meridional winds

Change in equatorial zonal currents and equatorial instability



- EUC is more realistic (stronger) in SCOAR.

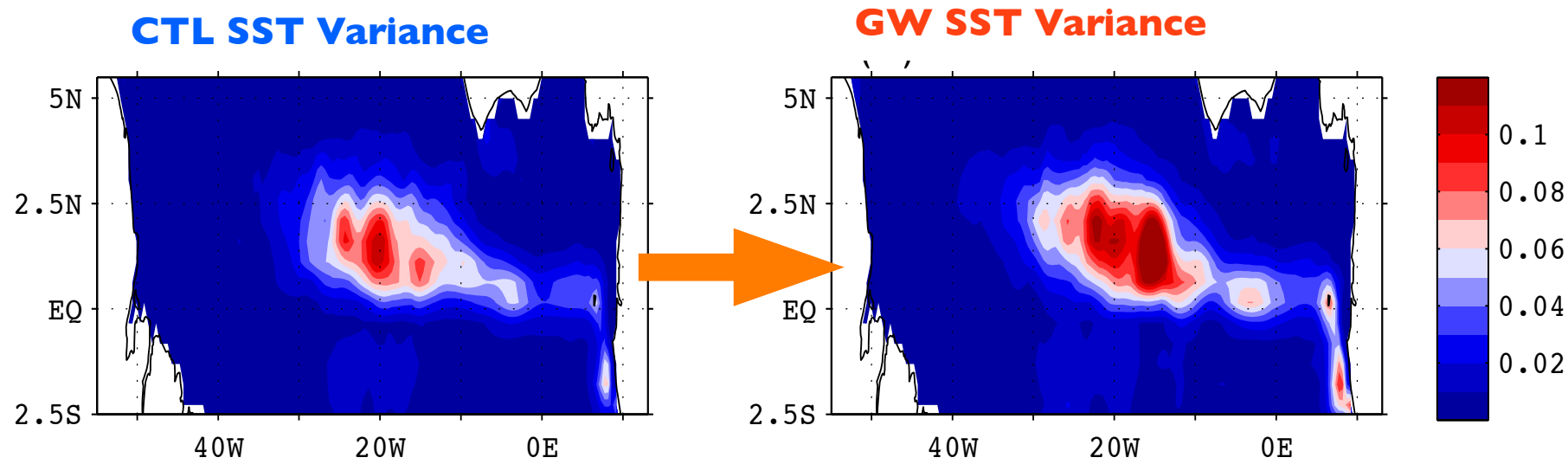
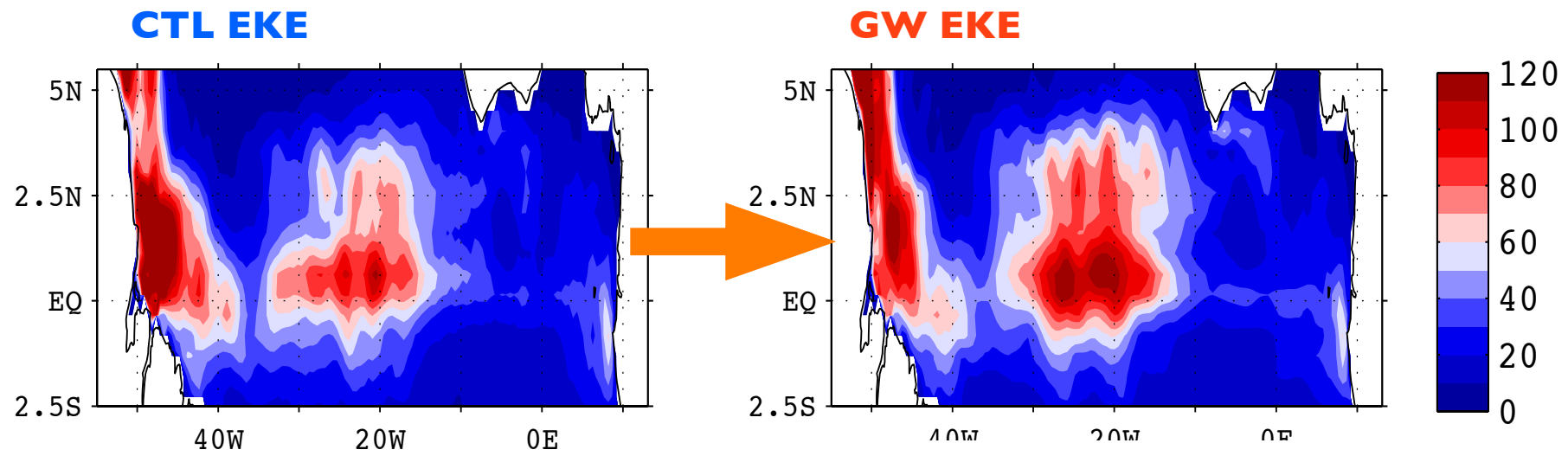
- Stronger cross-equatorial wind

➔ Stronger EUC
(Philander and Delecluse, 1983)

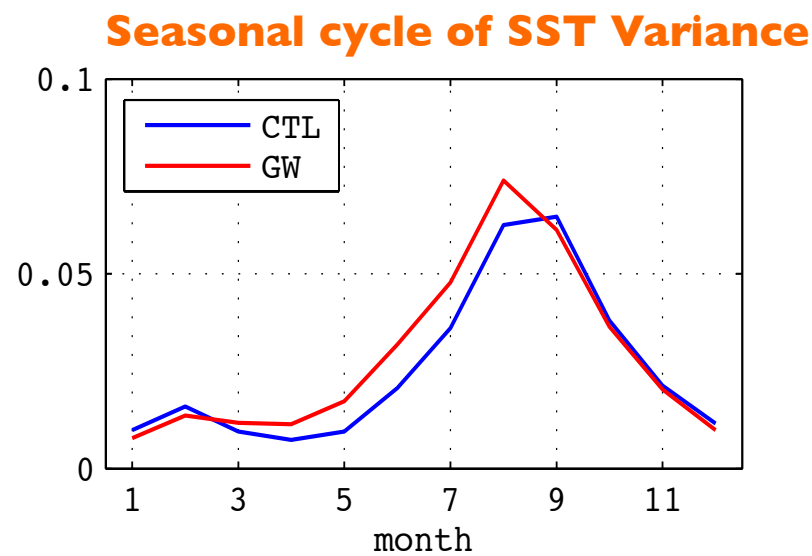
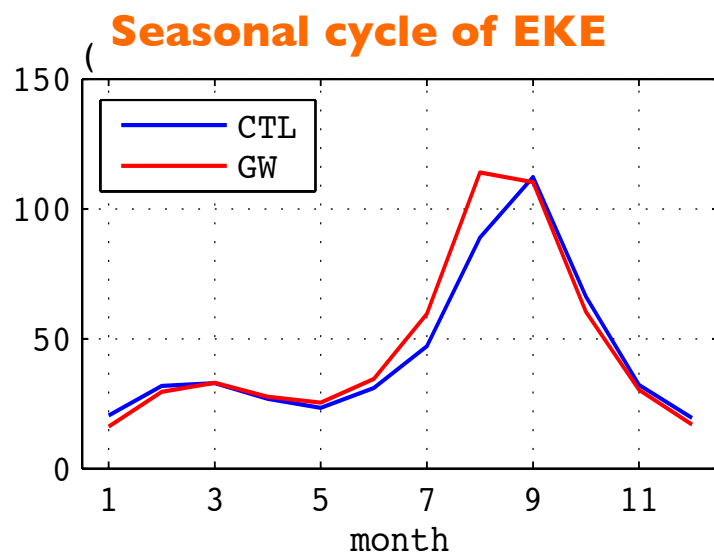
➔ Enhanced Barotropic and baroclinic instability

➔ Stronger TIWs

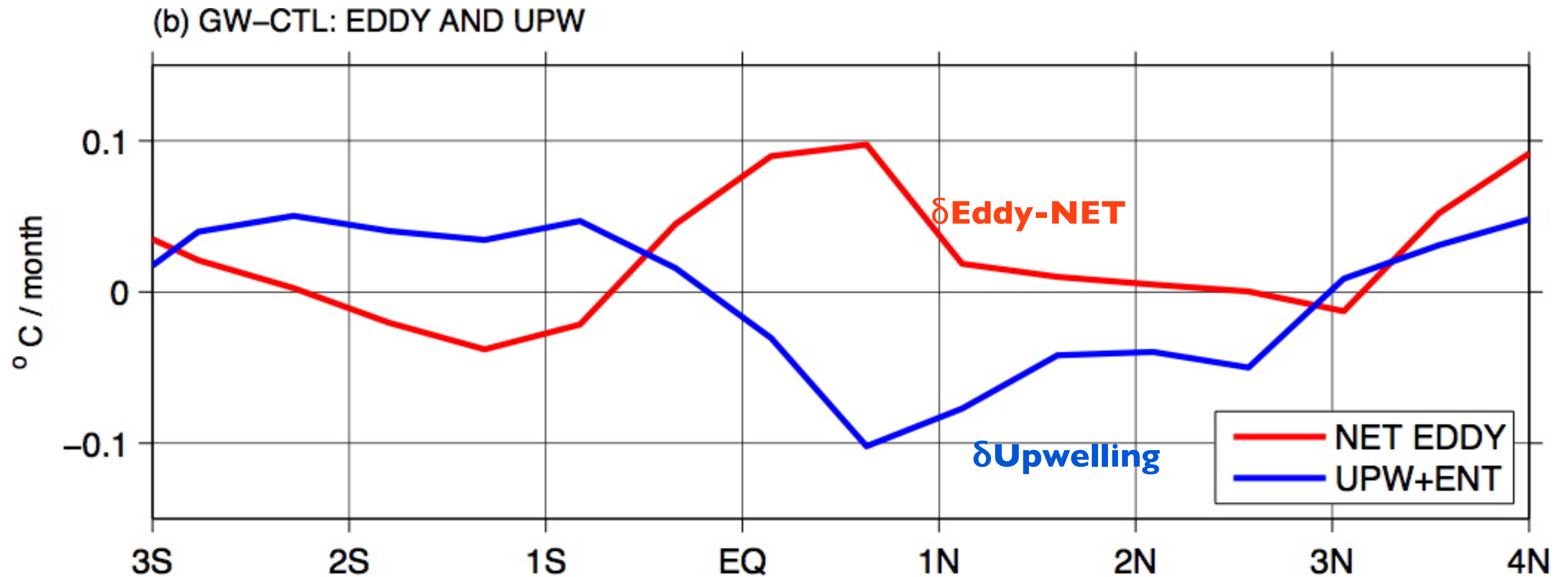
Strengthening of TIWs (20-40 day band-pass filtered EKE and SST variance)



- EKE and TIW-SST variance all become stronger during the cold season.



Annual mean mixed layer ocean heat budget (30W-10W)



- Equatorial upwelling (cooling) increases
 - Increased w' acting on climatological $dT/dz \gg$
Climatological $\langle w \rangle$ acting on dT'/dz due to radiative forcing.
- Net eddy heat flux (warming) increases, damping the effect of upwelling.

Conclusion and Discussion

- Downscaling is also important for study of oceanic role in weather and climate.
- Advantages: Better capture equatorial currents and mesoscale variabilities
- *Exploratory* research: Coupled downscaling of the IPCC climate change scenarios
- **Upwelling increases. TIWs increase. Impact the mean state.**
- Need to monitor TIW heat flux(zonal) for detection of warming signal.
- Need to resolve high-freq. processes in the model for global warming research.

Thanks!