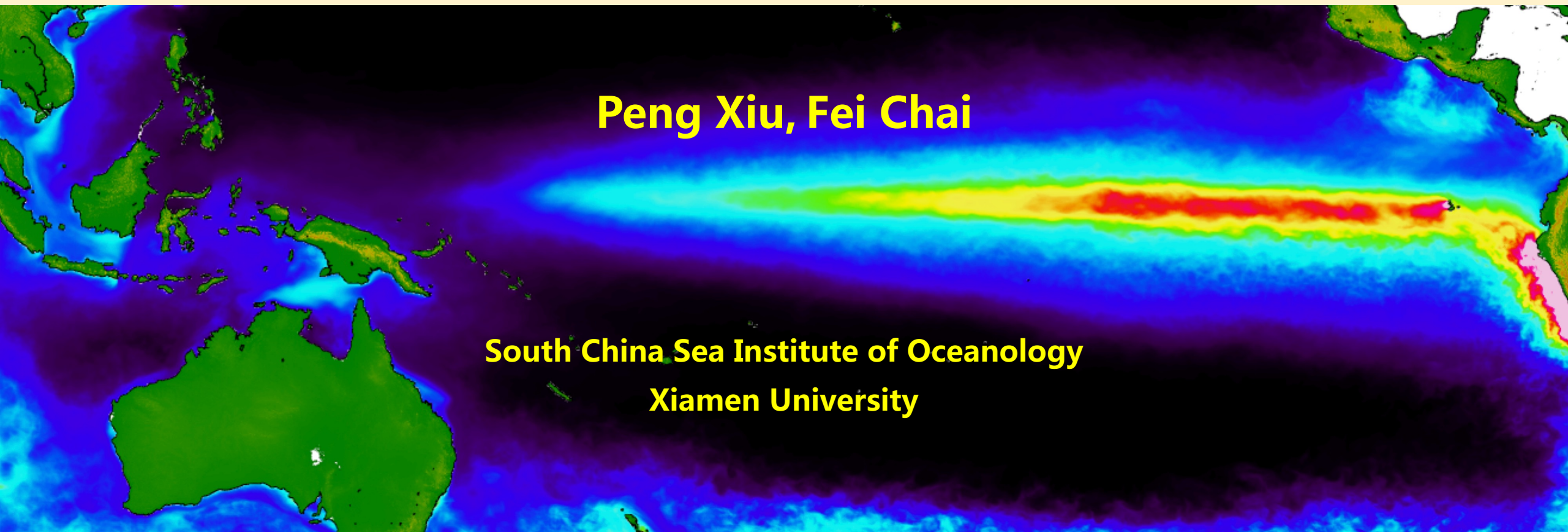


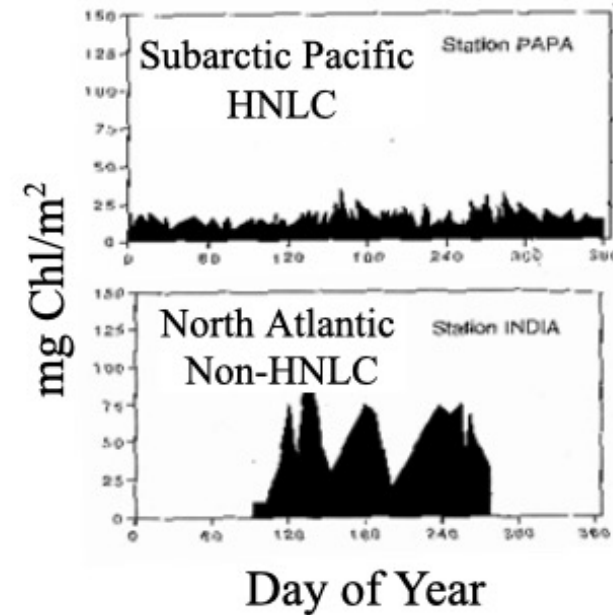
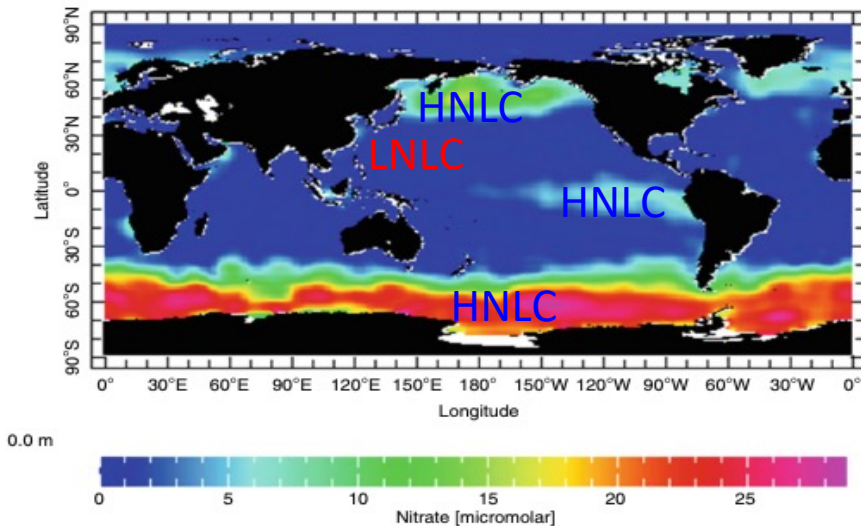
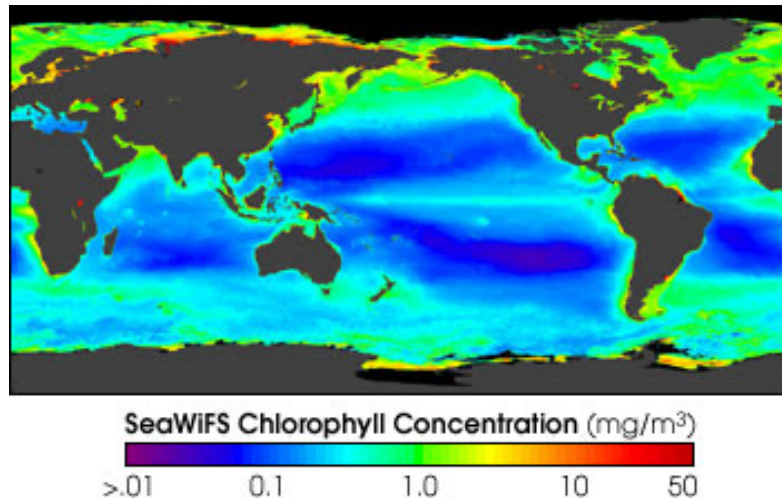
# Modeling the impact of iron on carbon cycling in the Pacific Ocean

**Peng Xiu, Fei Chai**

**South China Sea Institute of Oceanology  
Xiamen University**



# Phytoplankton chlorophyll and nitrate concentrations in the Pacific Ocean



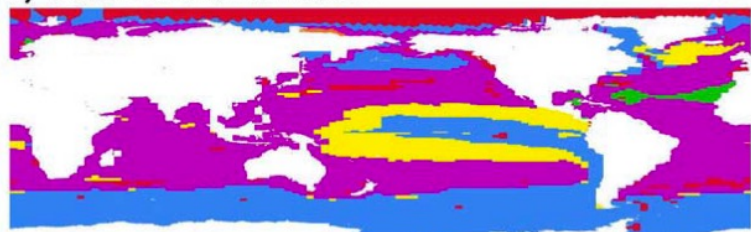
Frost, 1993; Parsons & Lalli, 1988

- About 30% of the ocean are with excess nutrients while lower phytoplankton chlorophyll than expected (High-Nitrate-Low-Chlorophyll regions; HNLC)
- Another ~30% of the ocean, mostly in subtropical gyres, are with low nutrients and low phytoplankton chlorophyll (Low-Nitrate-Low-Chlorophyll regions; LNLC)



# Limiting nutrient for phytoplankton growth in the North Pacific

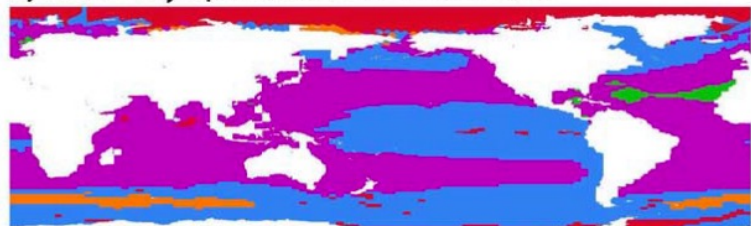
A) Diatom Growth Limitation



Nitrogen 55.73%, Iron 27.67%, Silica 12.54%, Phosphorus 1.405%  
Light 2.645%, Replete 0.000%

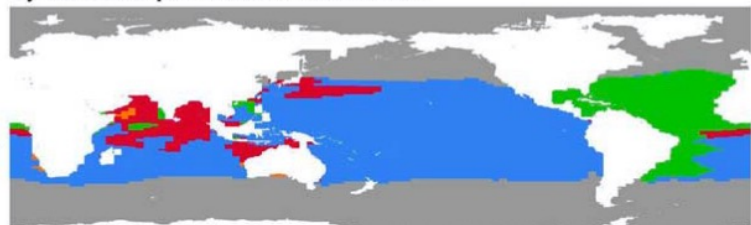
■ Nitrogen ■ Iron ■ Phosphorus ■ Silicon  
■ Light ■ Temperature ■ Replete

B) Small Phytoplankton Growth Limitation



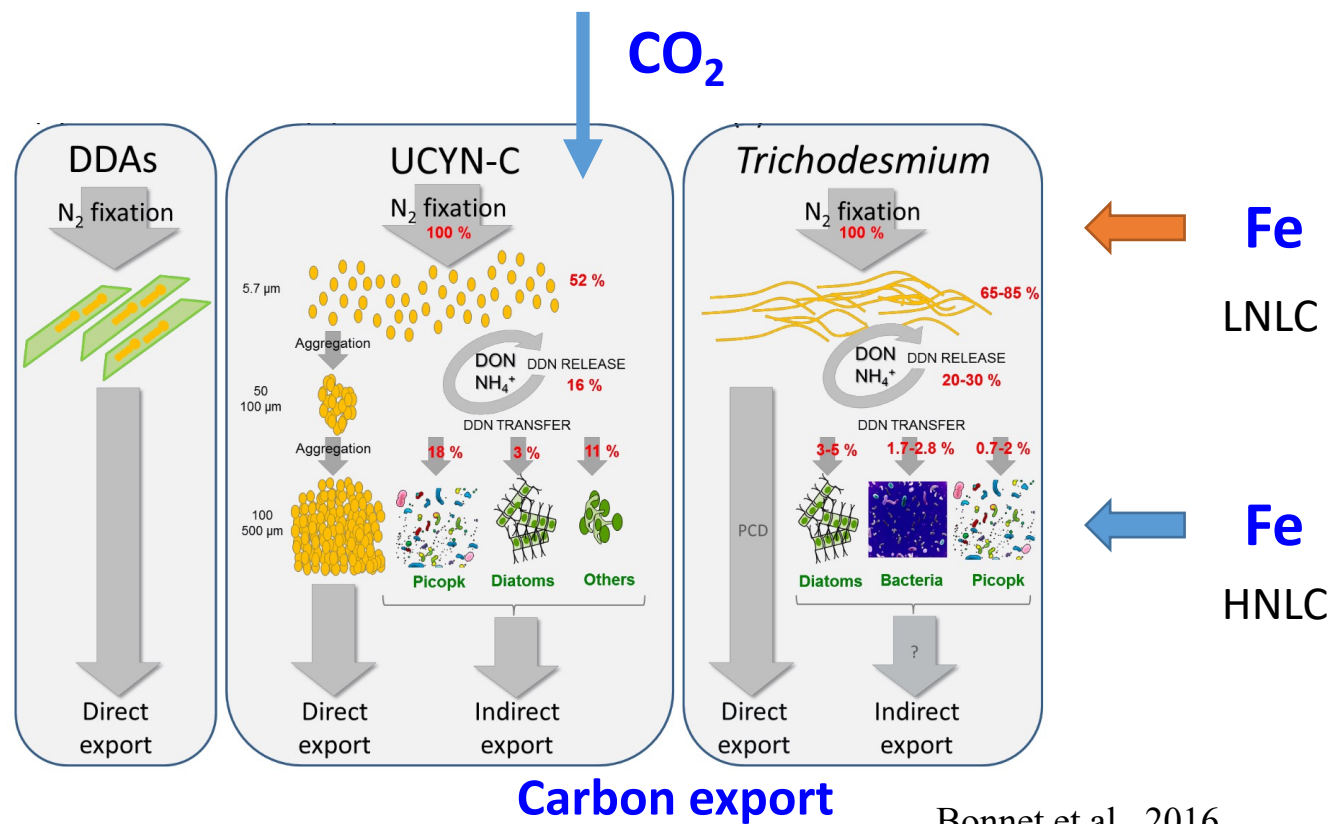
Nitrogen 55.88%, Iron 36.34%, Phosphorus 1.426%  
Light 3.788, Replete 2.556%

C) Diazotroph Growth Limitation



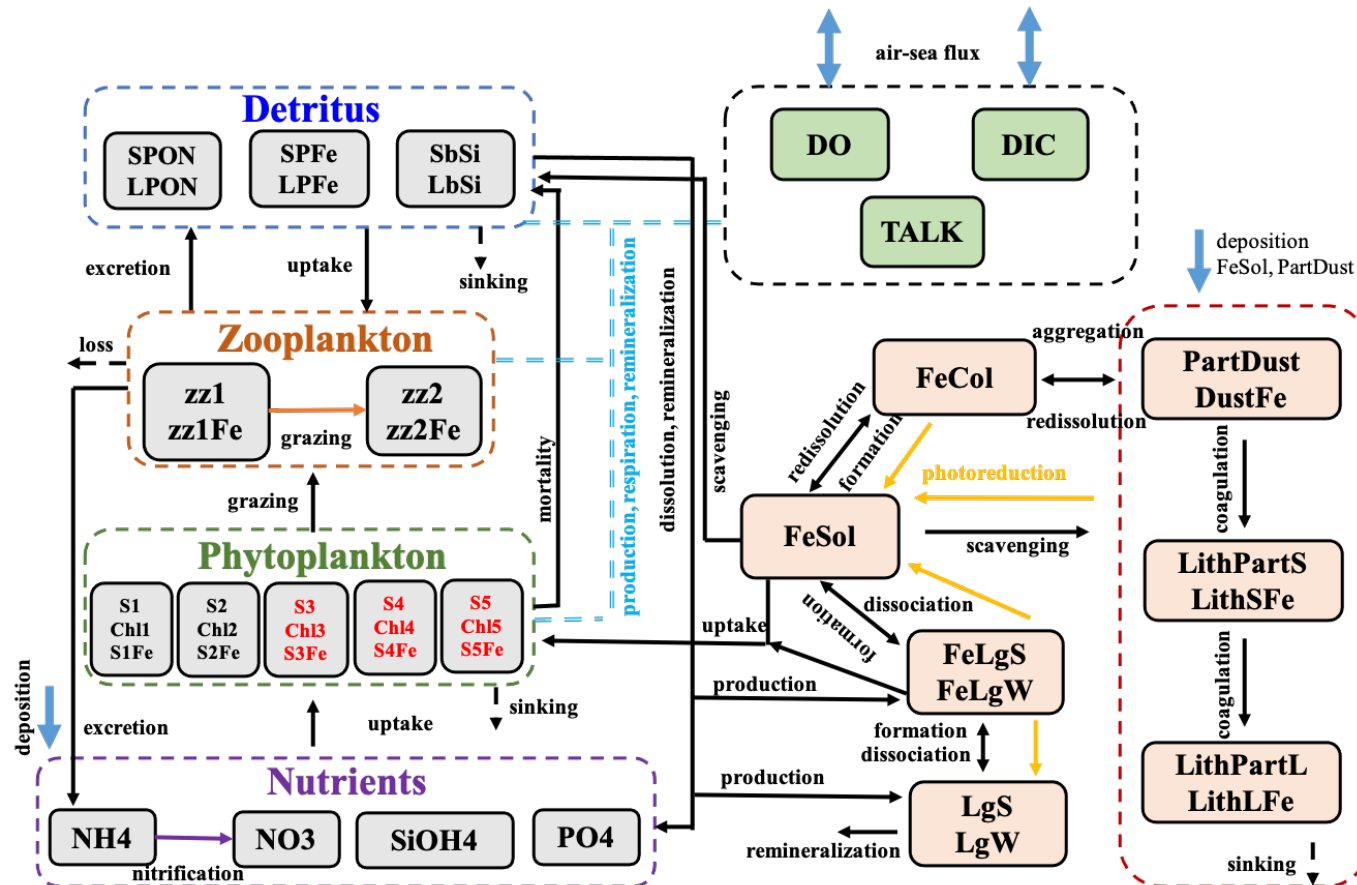
Nitrogen 0.000%, Iron 44.06%, Phosphorus 11.66%  
Light 7.072%, Temperature 36.81%, Replete 0.376%

Moore et al., 2004



- Iron is an essential nutrient for phytoplankton production across the North Pacific
- Variability in iron supply may lead to changes in the processes associated with biological carbon pump

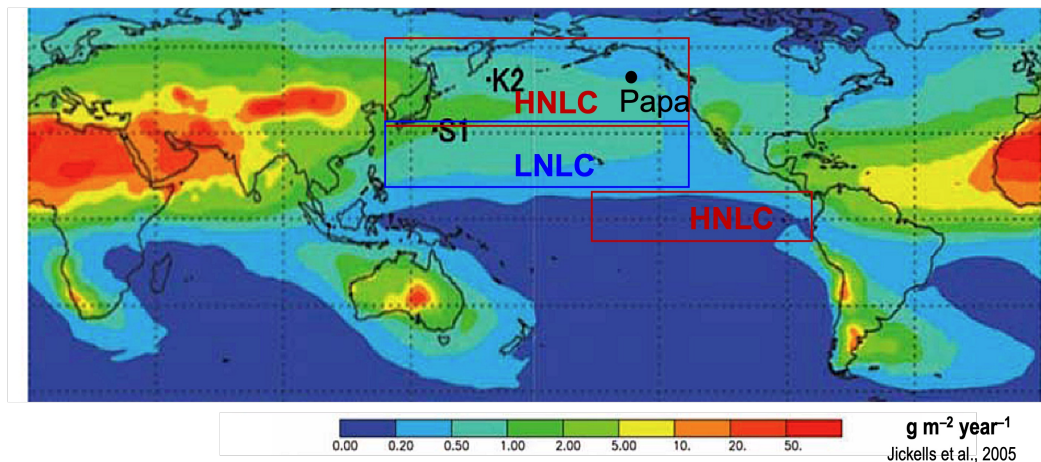
## CoSiNE-Fe



- Five phytoplankton groups including picoplankton, diatom, and three diazotrophs (unicellular cyanobacteria, Trichodesmium, diatom-diazotroph associations (DDA) )
- Iron cycles including soluble Fe, colloidal Fe, strong ligand Fe, weak ligand Fe, strong ligand, and weak ligand
- Atmospheric depositions of Fe, N, P, and lithogenic particles, and parameterized Fe sources from sediments and hydrothermal vents
- Three size classes of particles with different parameterizations
- A new light attenuation scheme with a dual-band model. Phytoplankton photoacclimation was parameterized



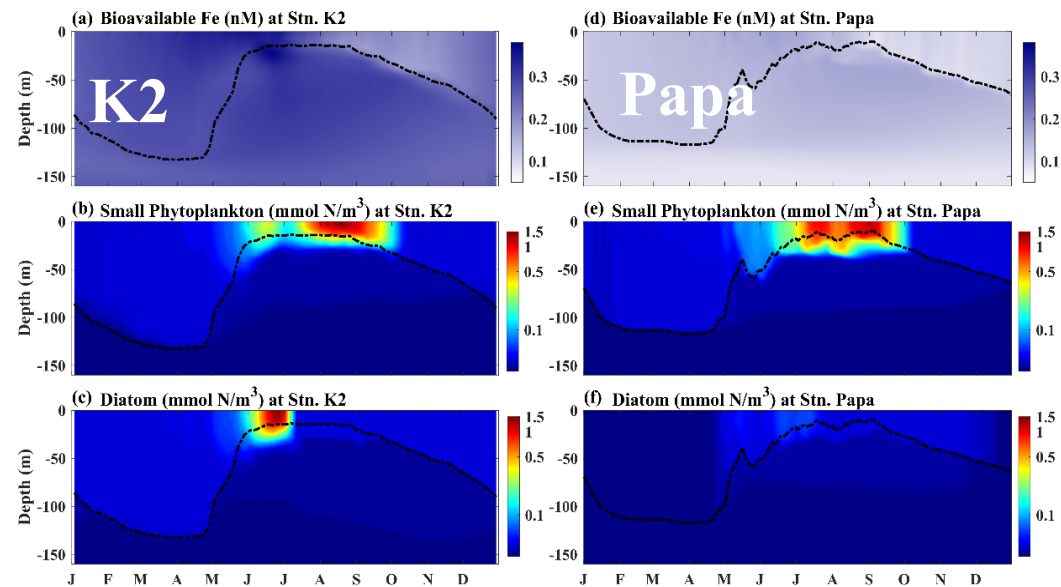
# One-dimensional ROMS-CoSiNE-Fe model



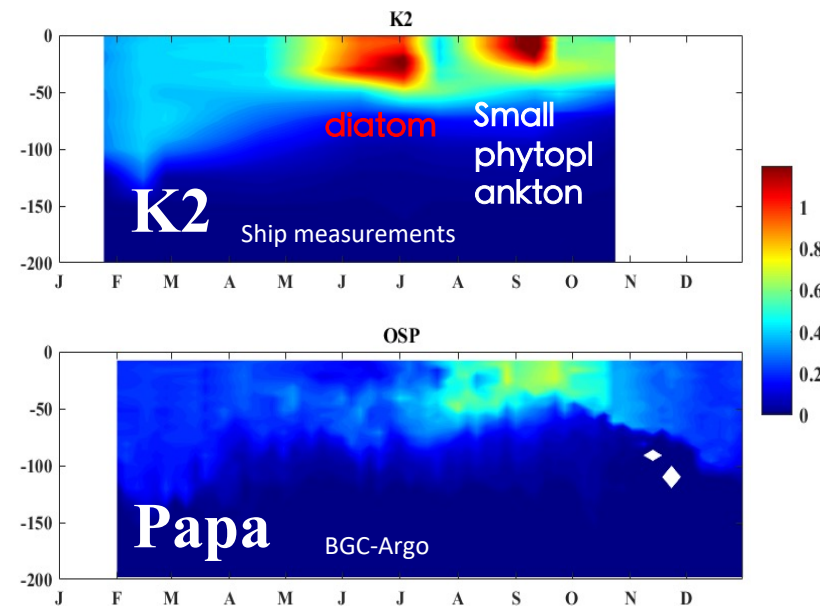
- Bioavailable Fe is higher in the western region (K2)
- Small phytoplankton peaks in late summer in both regions, with a relatively larger magnitude in the western region (K2) than the eastern region (Papa)
- Diatom in the the western region (K2) peaks from late spring to summer, while is constantly low in the eastern region (Papa)

The west-east gradient in dust Fe deposition shapes the difference of marine ecosystems between the western and eastern subarctic Pacific Ocean

## Model

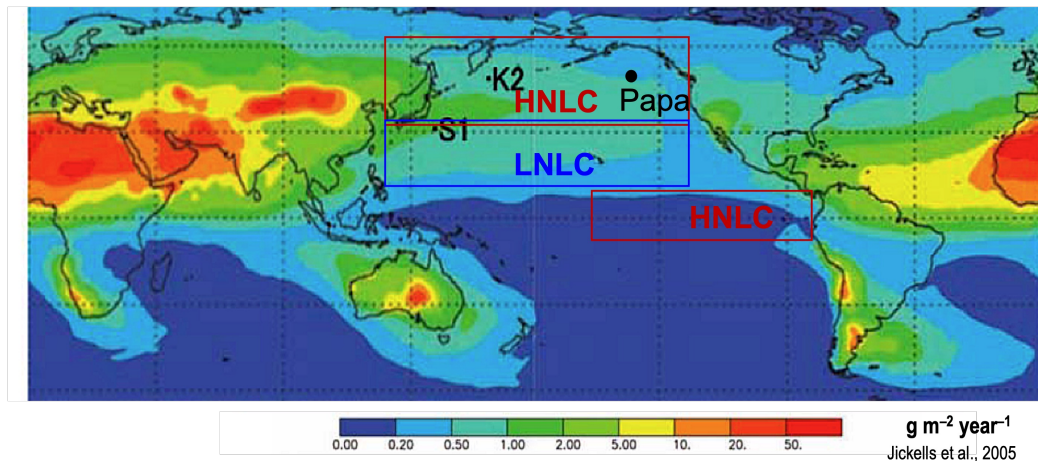


## Observations

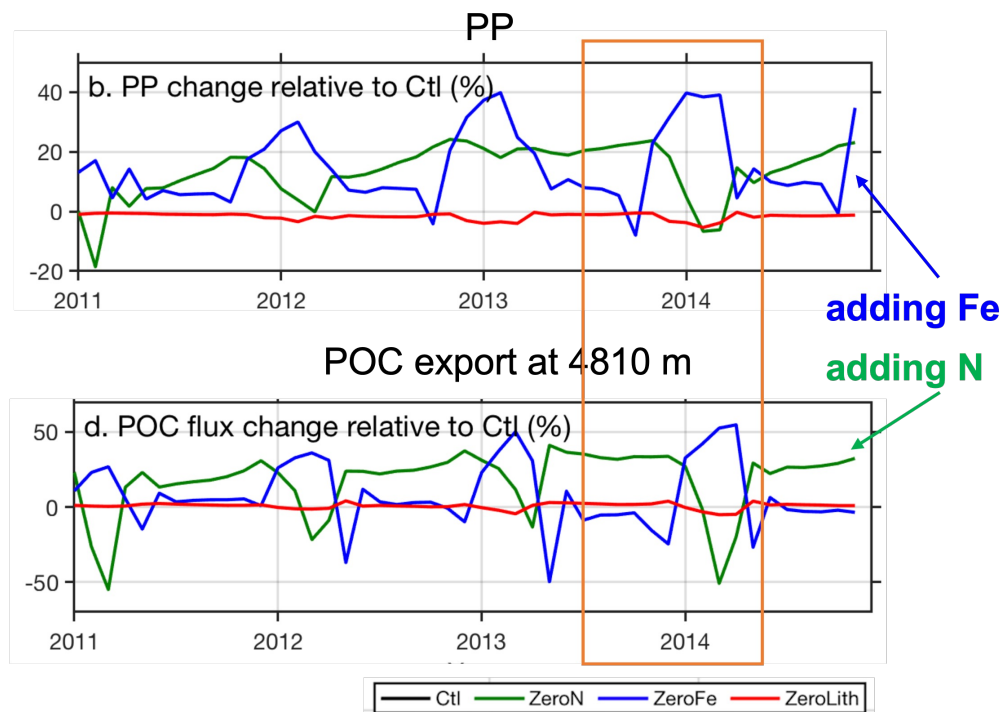


Zhang et al., 2021 FMS  
Zhang et al., 2021 MEPS

# One-dimensional ROMS-CoSiNE-Fe model



- S1 (145°E, 30°N) is in the NPSG, but experiences strong winter mixing
- Phytoplankton is limited by nitrate in summer and fall, and nitrogen fixation is not strong in this region
- Atmospheric Fe deposition is still a significant input, especially from late winter to early spring
- A moored sediment trap at 200m, 500m, 4810m (Honda et al., 2002) from JAMSTEC

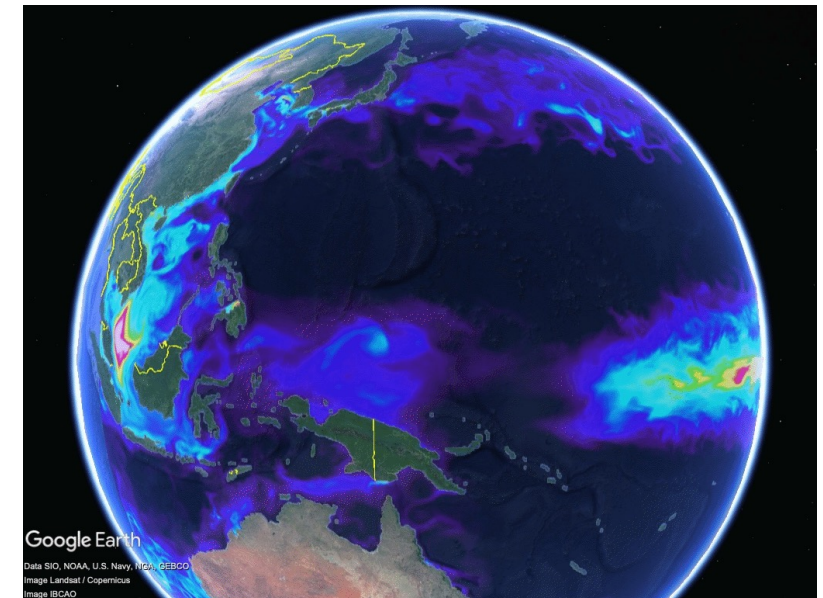
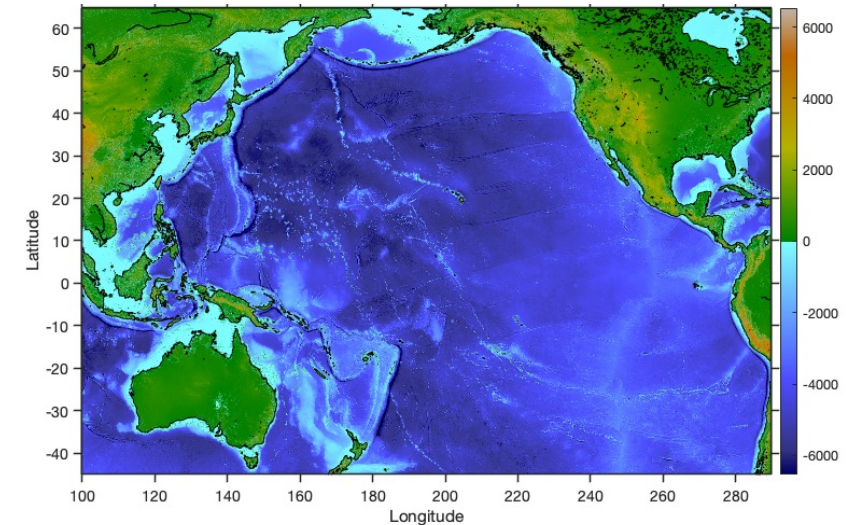


- Atmospheric deposition of N and Fe **increases 0-200m PP**
- Atmospheric **N** stimulates picoplankton growth in summer and fall with increased grazing pressure on winter diatoms and consequently **reduces deep ocean export in winter**
- Atmospheric **Fe** deposition matters in winter and **stimulates diatoms growth and deep ocean export**

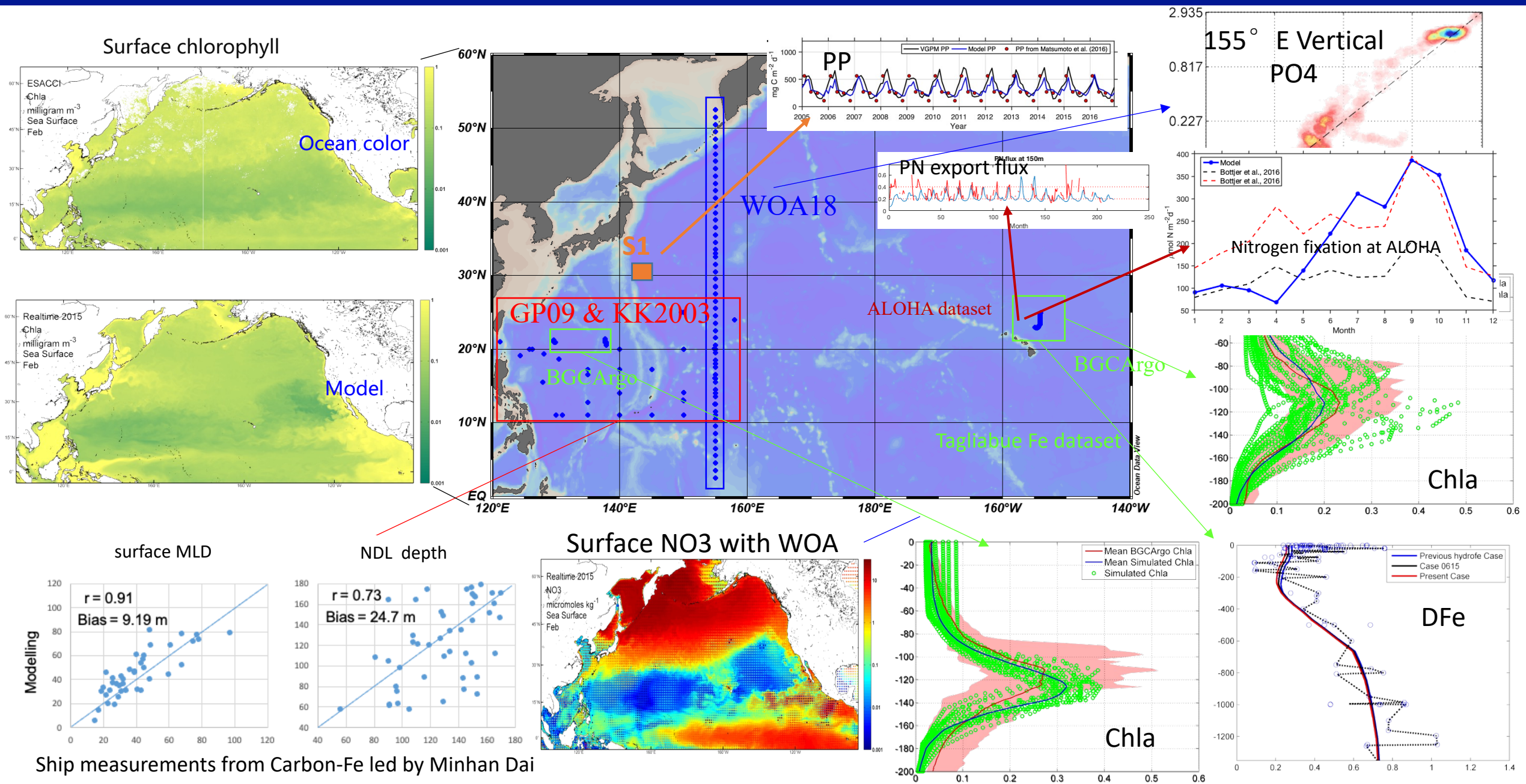


## Pacific ROMS-CoSiNE-Fe

- ROMS model: 99~289.9°E; -44~64.7°N
- **Horizontal resolution**: 1/8° (~10 km)
- **Vertical layers**: 60
- Topography data: SRTM30plus
- Lateral boundaries: HYCOM, WOA for biogeochemical variables
- Forcings: ERA5
- Rivers: Climatological monthly river discharge
- Tides: TPXO7.2 from OSU (M2, N2, S2, K2, K1, O1, P1, Q1)
- Vertical mixing: MY level-2.5 closure
- **Atmospheric deposition**: Monthly data from Chien et al. (2016)
- **Hydrothermal vent Fe flux**: Tagliabue et al. (2010)
- **Sediment Fe flux**: A function of depth (Moore et al., 2004; Aumont et al., 2015)



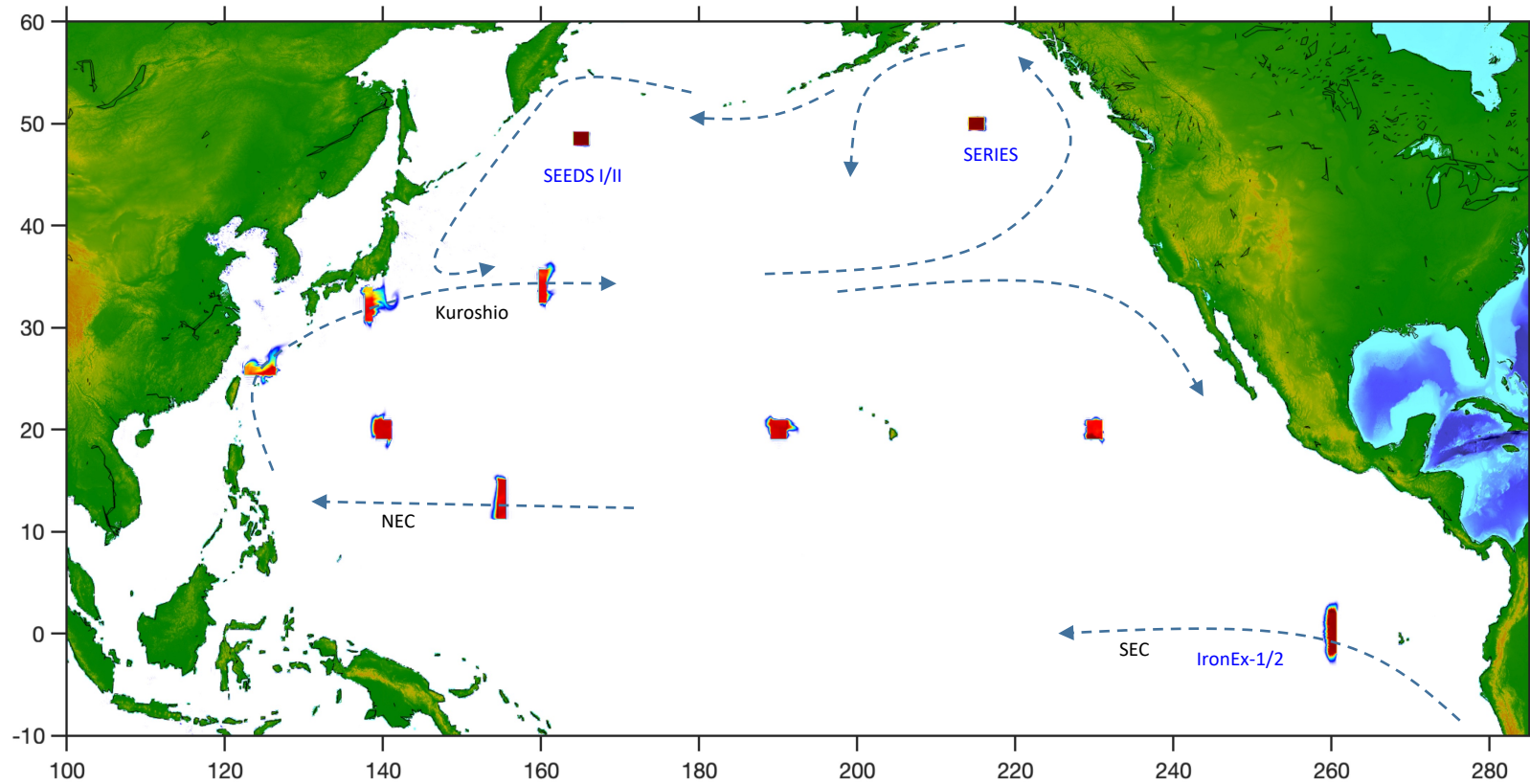
## Model validations: Satellite derived products, BGC-Argo, WOA, in situ measurements





# Modeling OIF experiments

## Ten locations in the Pacific Ocean



- **Season**

- Two cases: starting from **April 1** and from **July 1 in 2017**
- Both cases run **3 months**

- **Size**

- For barges: **200 km long x 12.5 km wide, 400 km long x 100 km wide**
- For boxes: **50 km x 50 km, 200 km x 200 km**

- **Depth**

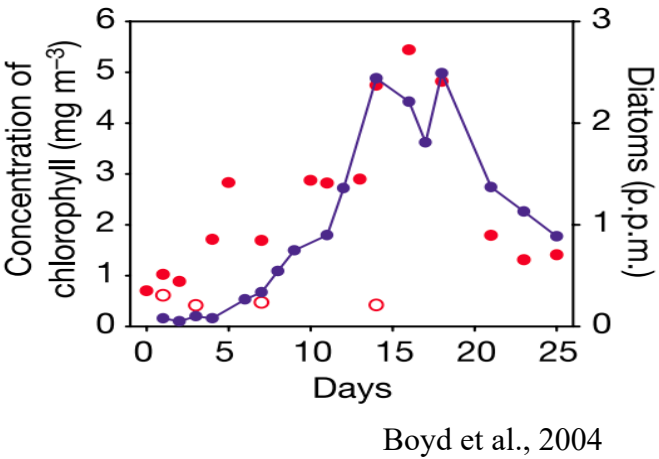
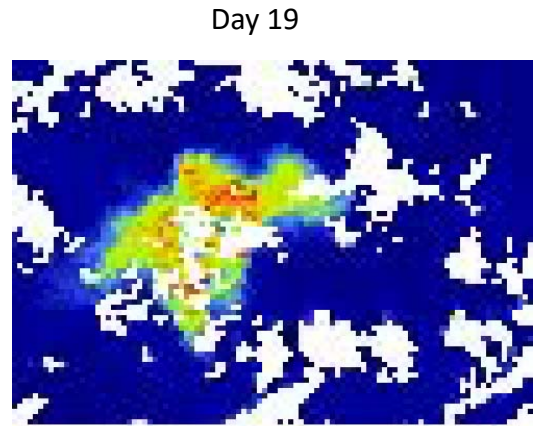
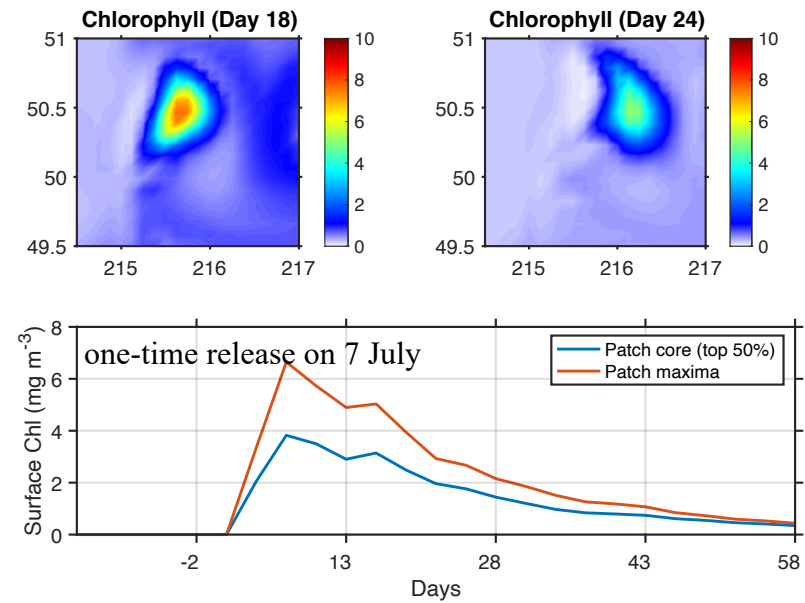
- In the upper **10 m**

- **Iron amount**

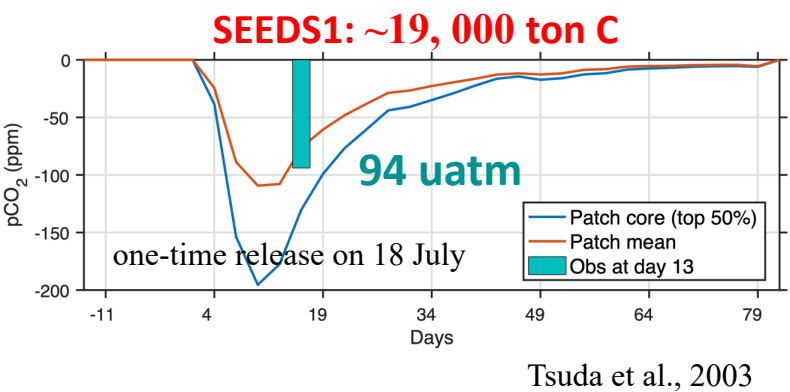
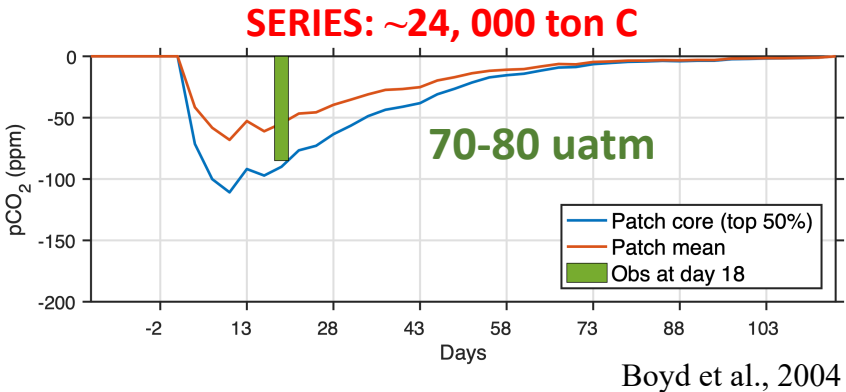
- Increase DFe concentration to **1.0 nM**
- Continuous release for **15 days**

# Modeling OIF experiments

- Comparison with data during SERIES



- Surface  $\text{pCO}_2$  change



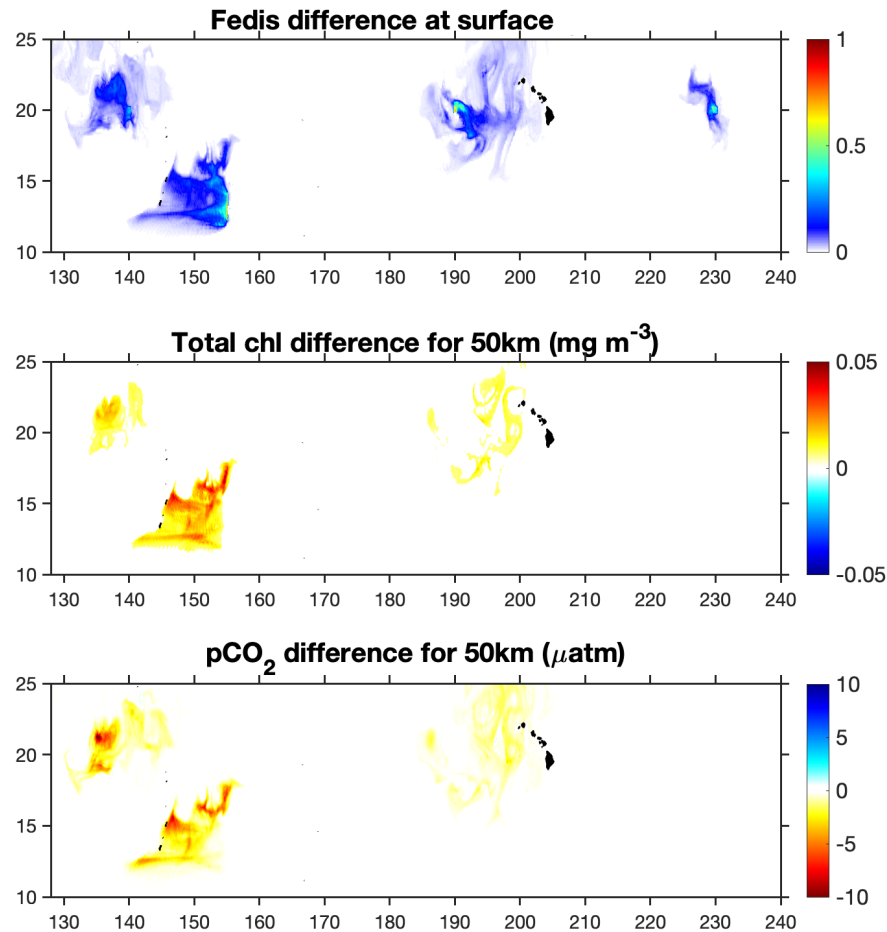
- Model results are generally consistent with previous OIF experiments in subarctic region
- Adding Fe can drive surface  $\text{pCO}_2$  drawdown
- The model can track the fertilized patch over its lifetime and estimate the integrated effect



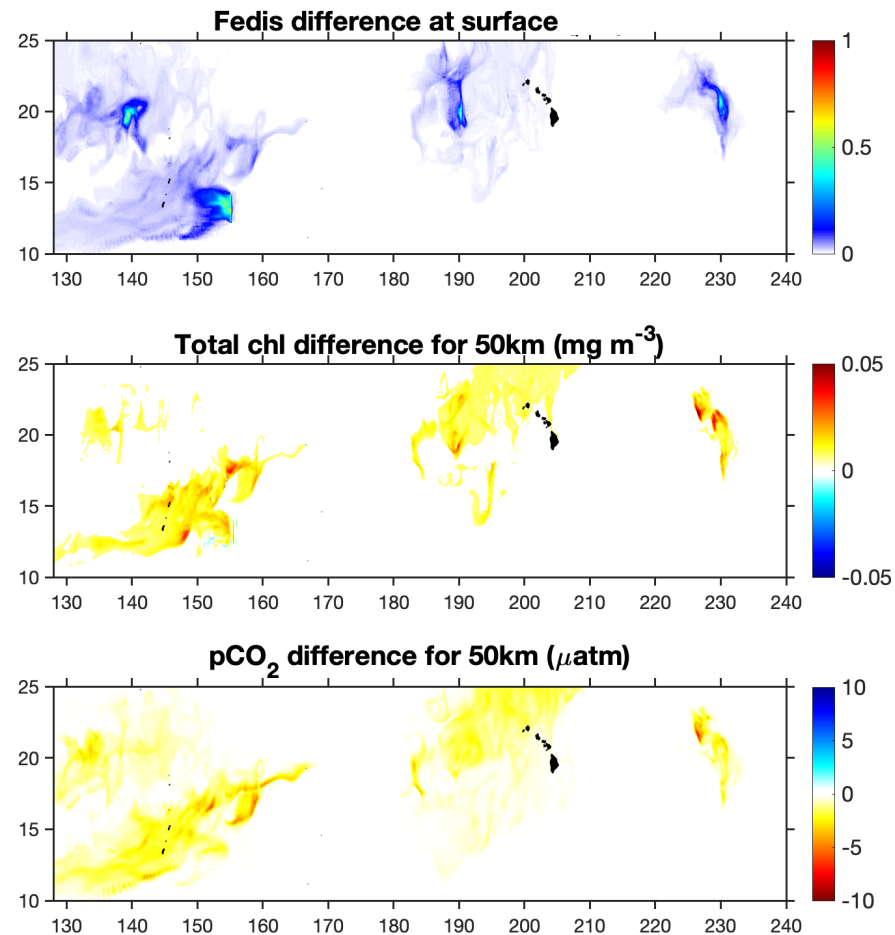
# Modeling OIF experiments

- Surface response averaged in three months

**Spring**



**Summer**

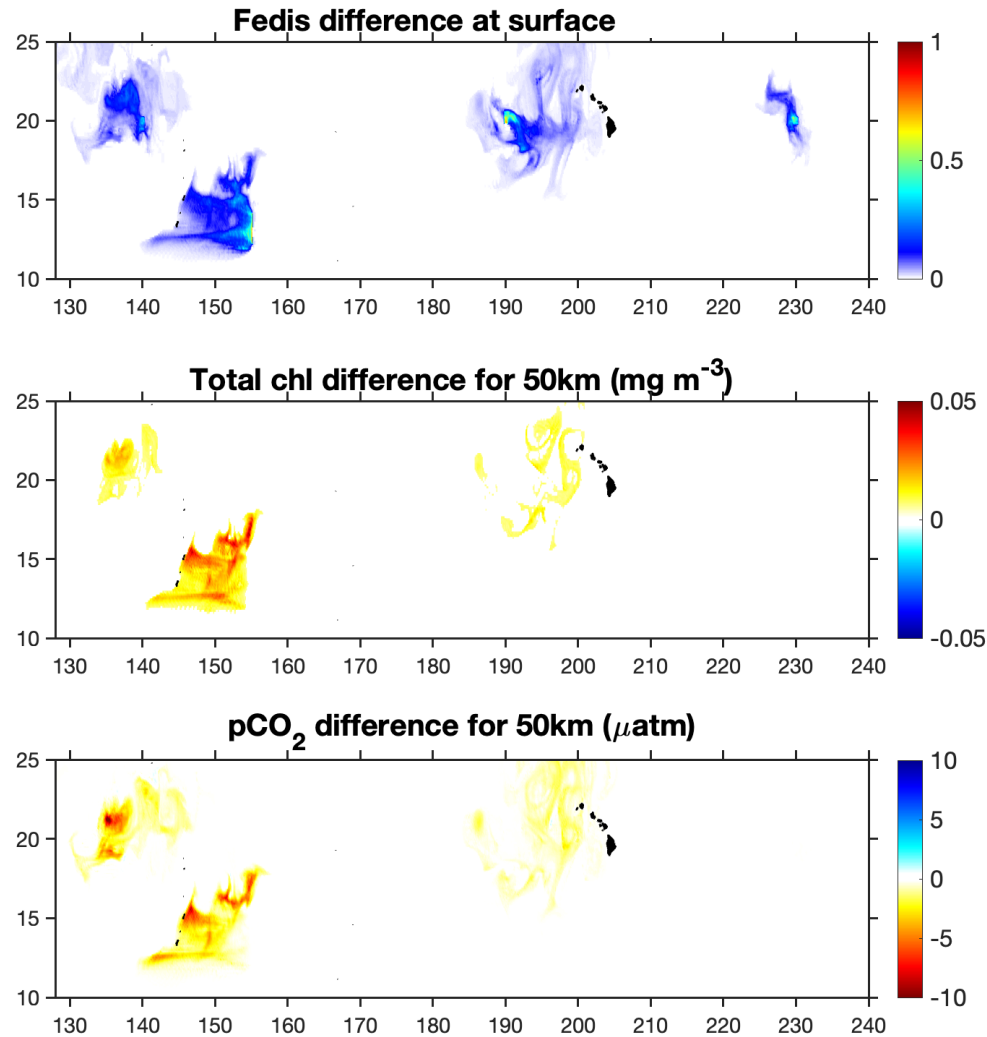


- The eastern patch shows no biological response in spring, but shows response in summer
- Phyto response is relatively stronger in the western gyre than in the central gyre in spring
- During three months, the summer experiment covers more space than the spring experiment; the spring experiment is localized with high values

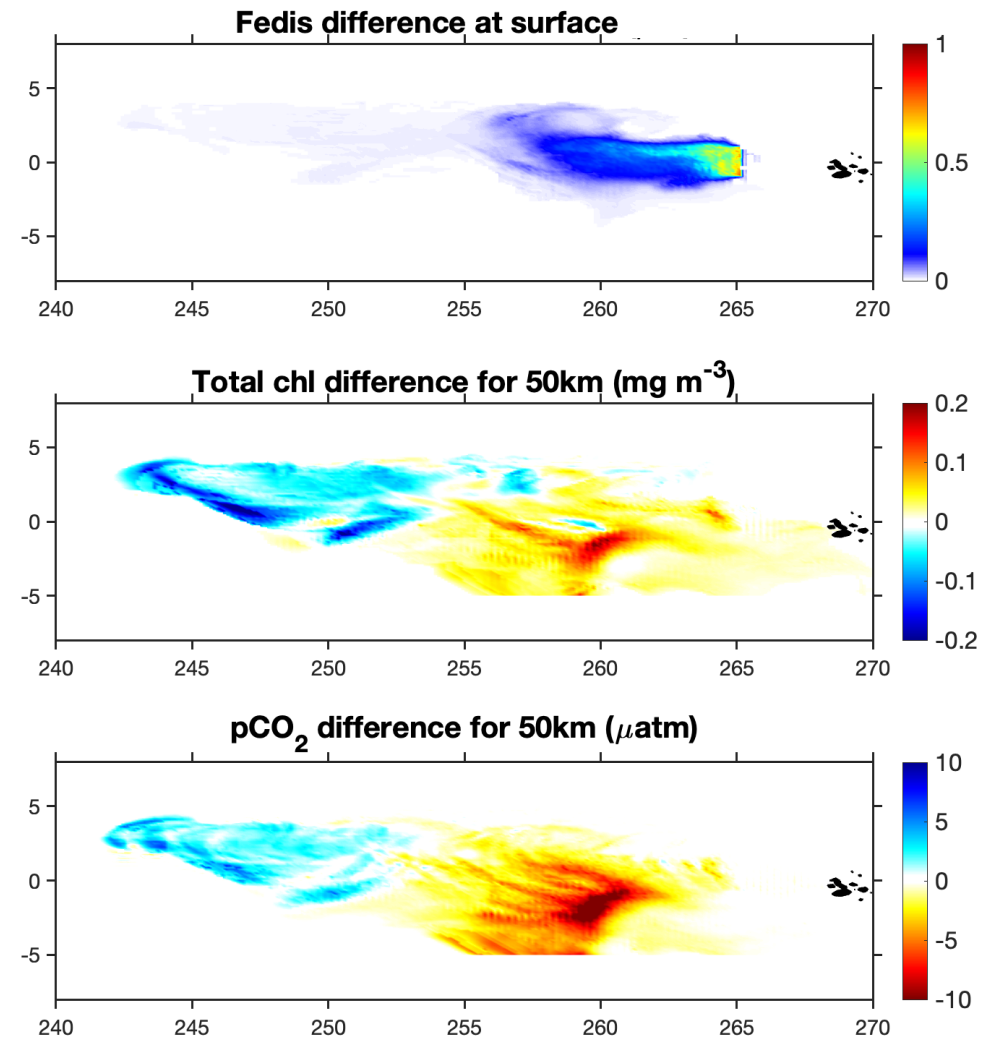
# Modeling OIF experiments

- Surface response averaged in three months (April-June)

**NPSG**



**EEP**

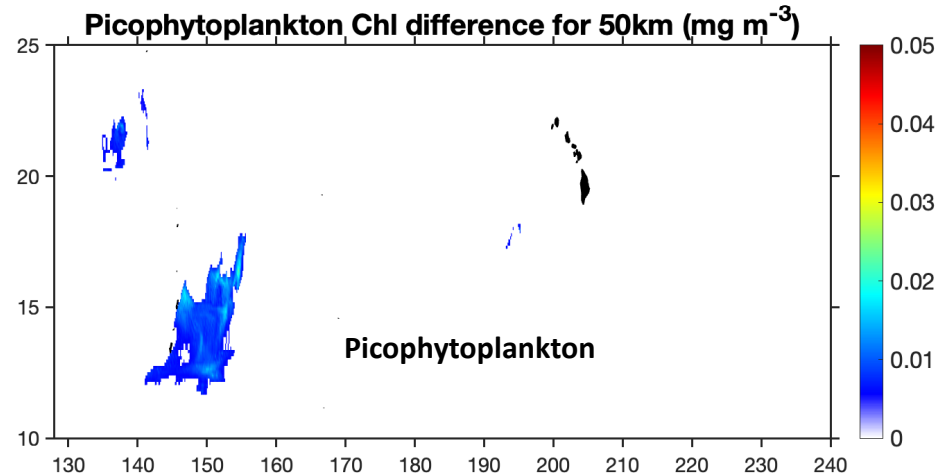




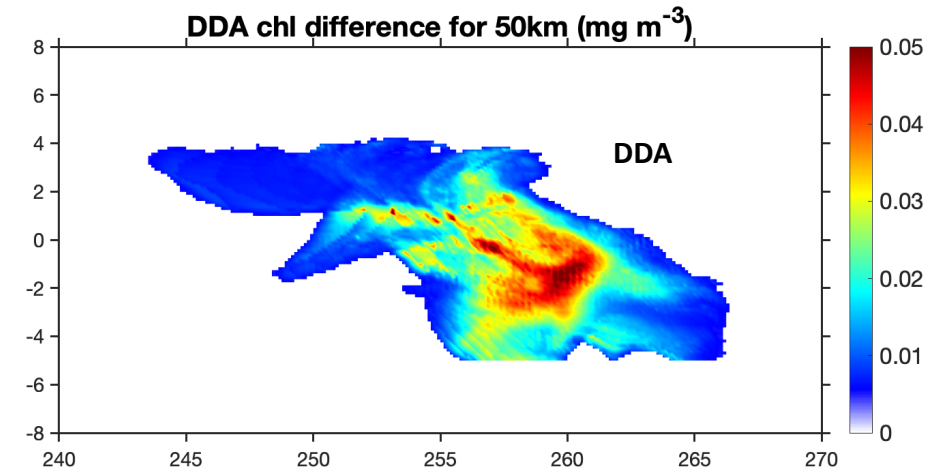
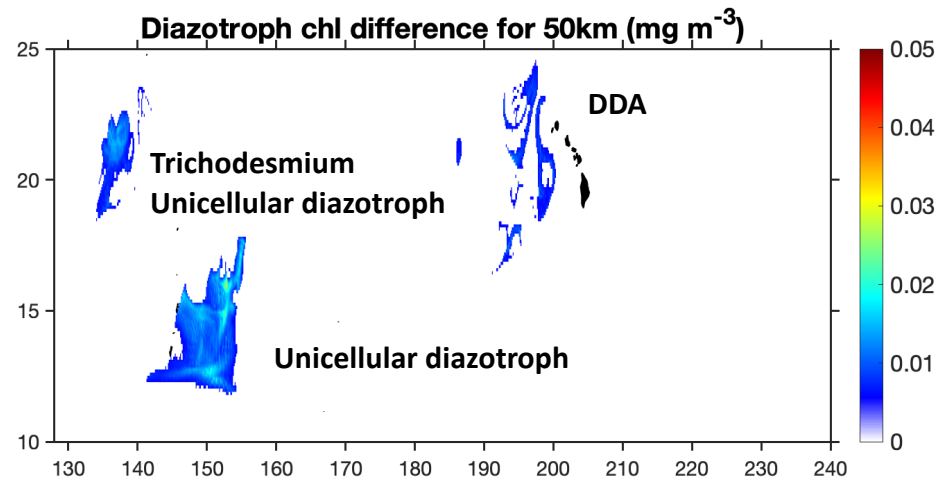
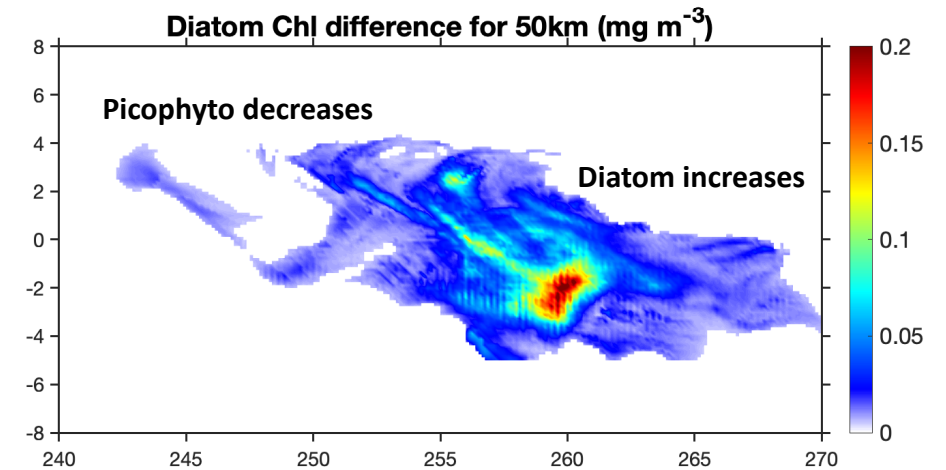
# Modeling OIF experiments

- Surface response averaged in three months (April-June) : Related to different phytoplankton species

## NPSG

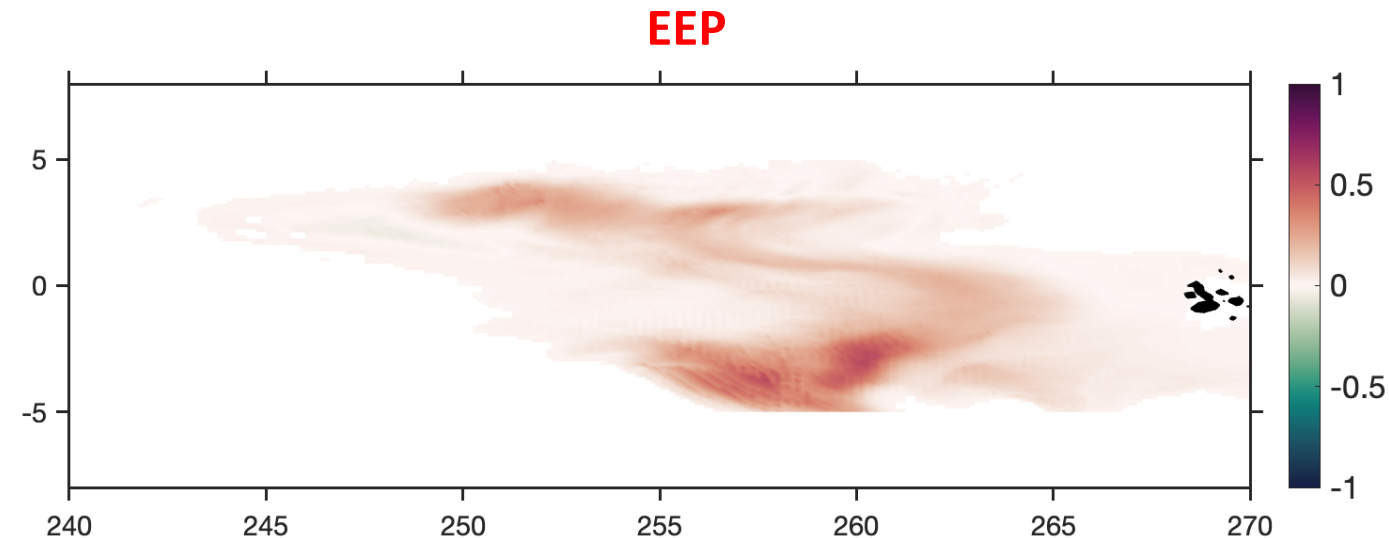
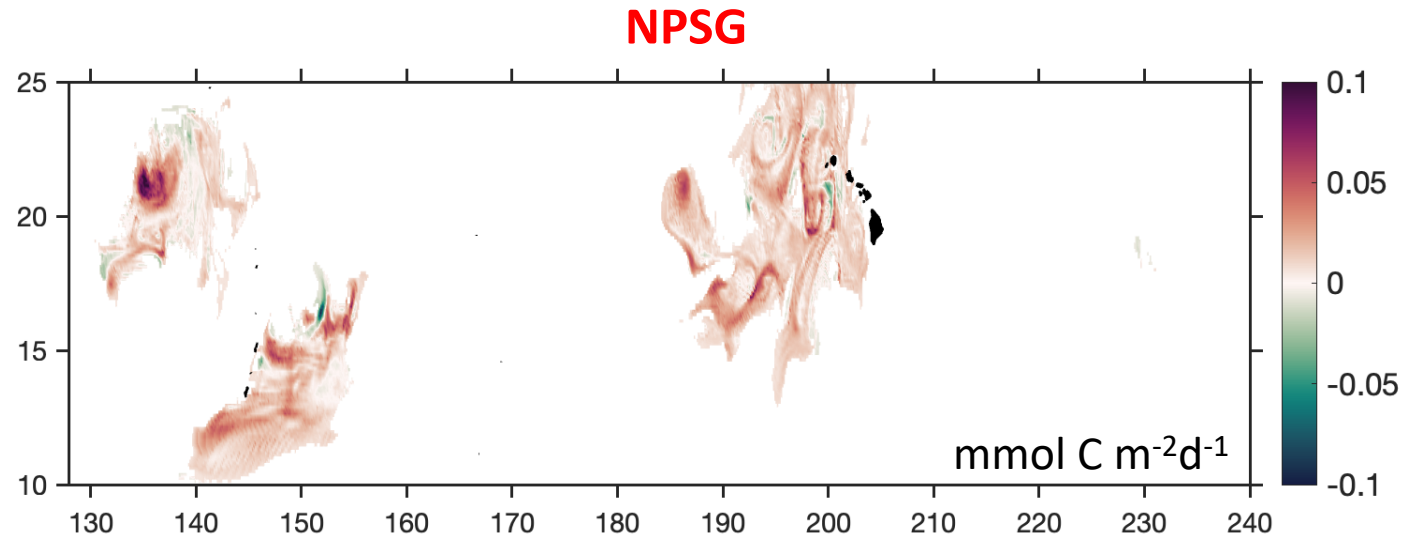


## EEP

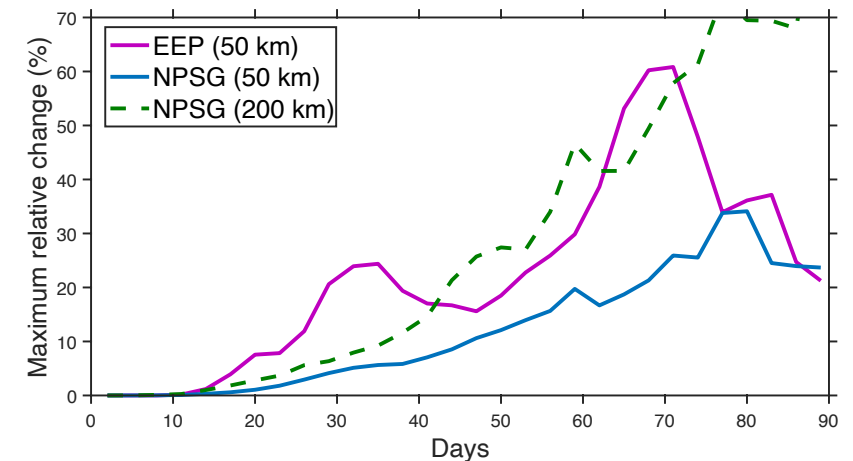


# Modeling OIF experiments

- POC export flux at 200 m averaged in three months (April-June)



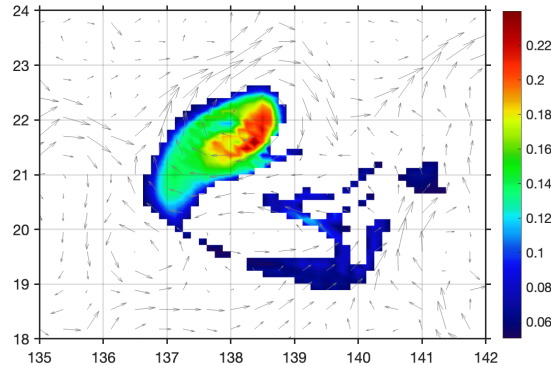
- Both NPSG and EEP show changes in the POC export flux at 200 m
- The magnitude is larger in EEP than in NPSG
- The downstream change in EEP is not seen in the POC flux at 200 m, although it changes surface  $\text{pCO}_2$



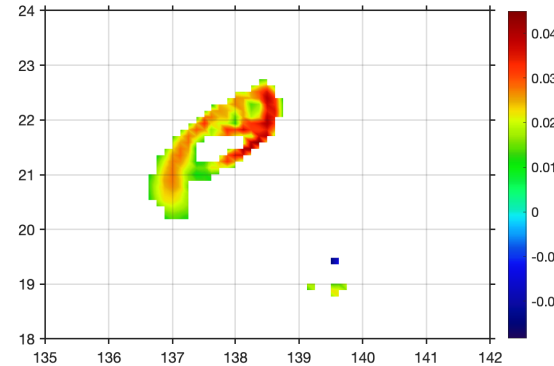
# Modeling OIF experiments

- Spatial distributions of patches on day 45

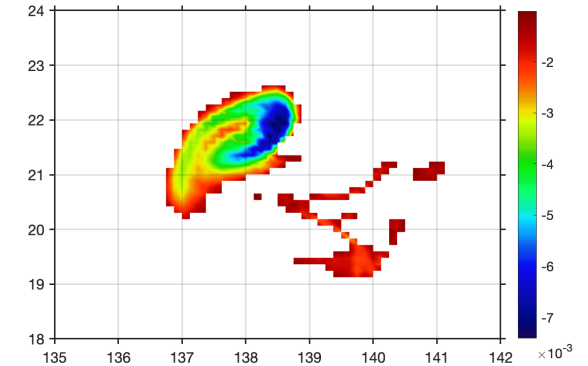
Fe patch



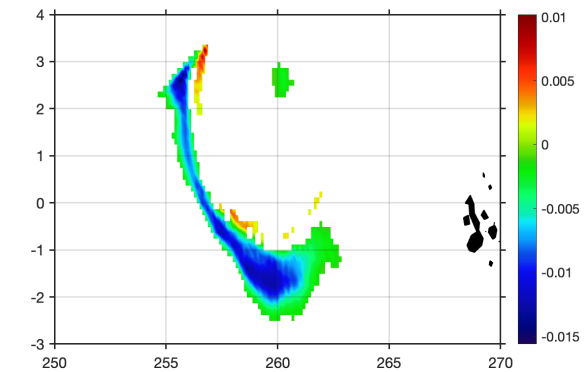
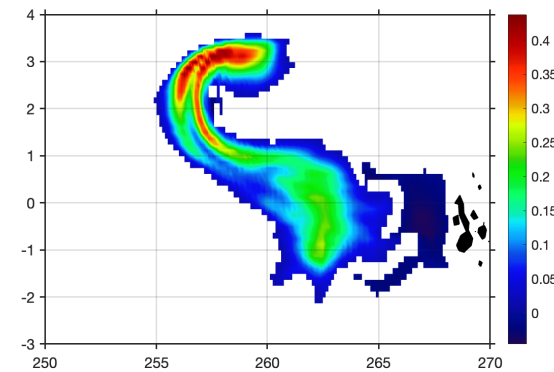
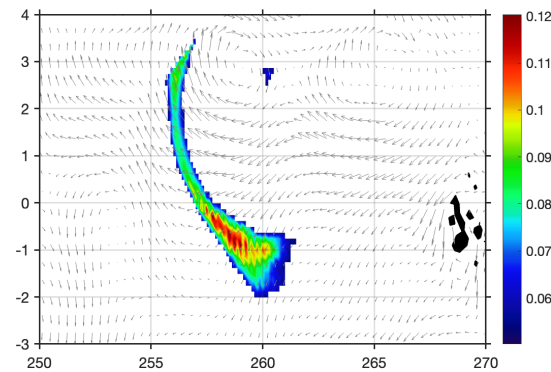
POC flux patch



Air-sea CO<sub>2</sub> flux patch



EEP



- In the western gyre, the Fe patch is in a similar shape as the POC flux patch and the air-sea CO<sub>2</sub> flux patch, regulated by a mesoscale eddy
- In the EEP, the shape of the Fe patch is similar to the air-sea CO<sub>2</sub> flux patch, but different from the POC flux patch



# Summary

- A new biogeochemical model (CoSiNE-Fe) was built and coupled with the high-resolution ROMS model for the Pacific Ocean. The coupled model was validated against satellite data, BGC-Argo data and in situ measurements
- OIF experiments were simulated and evaluated by the coupled model in different regions and different seasons in the North Pacific Ocean
- In the western gyre, adding Fe may stimulate diazotroph (Trichodesmium and Unicell) and picophytoplankton growth, which can draw surface  $p\text{CO}_2$  and affect air-sea  $\text{CO}_2$  flux. In the EEP, diatoms are simulated, which contributes to the POC export.
- The efficiency of OIF in changing air-sea  $\text{CO}_2$  flux is comparable between the western gyre and the EEP, but is lower in exporting POC in the western gyre than in the EEP
- The Fe patch at the surface may be in different shapes and locations from the POC flux patch at 200 m (EEP)